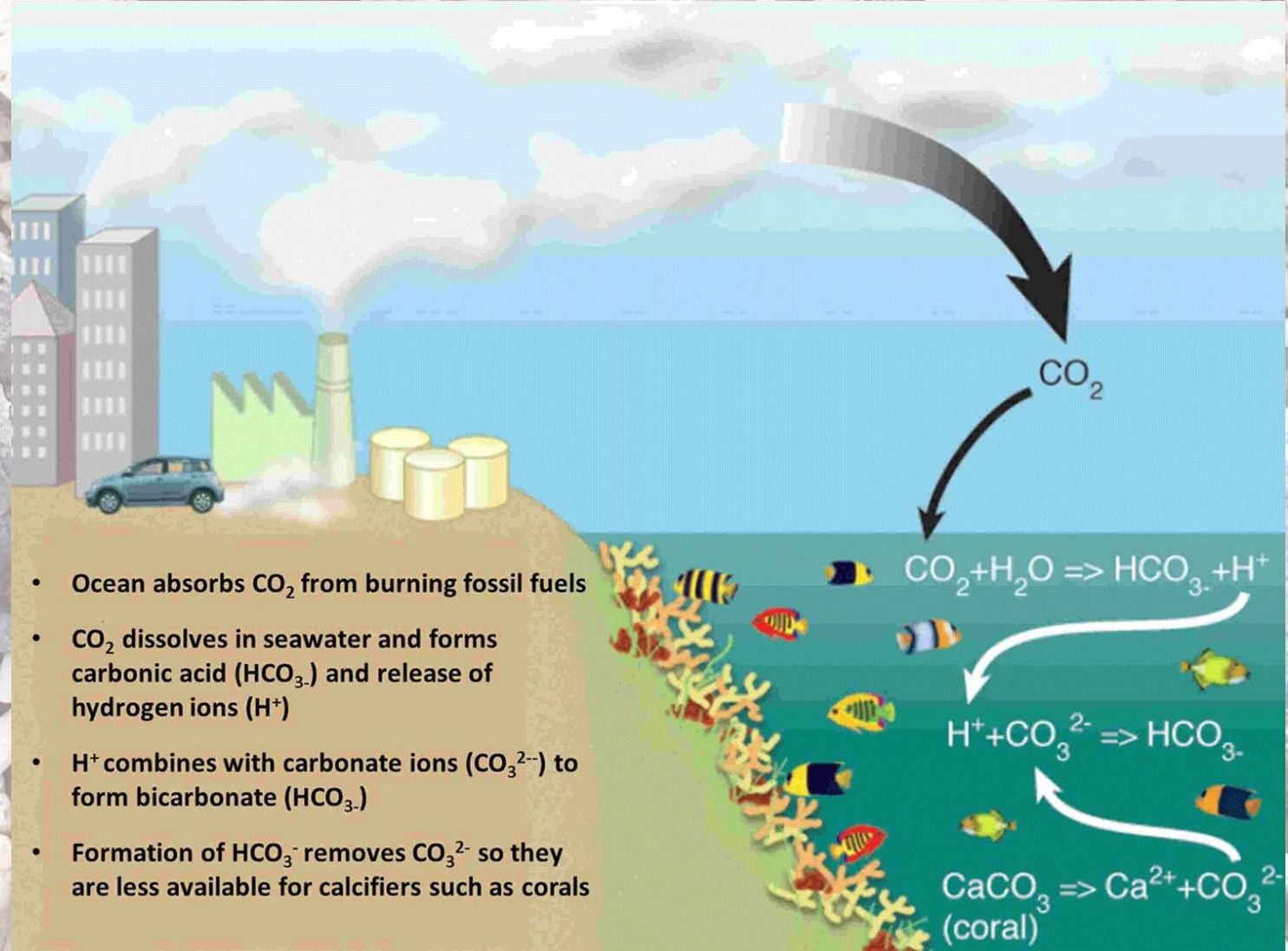


WHAT IS OA?



Summary of OA Science

- Well documented, **progressive increase in the acidity** of the ocean
- Acidity *increases* as the pH *decreases*
- OA is also changing seawater carbonate chemistry. The concentrations of *dissolved CO₂*, hydrogen ions, and bicarbonate ions are increasing, and the concentration of carbonate ions is decreasing

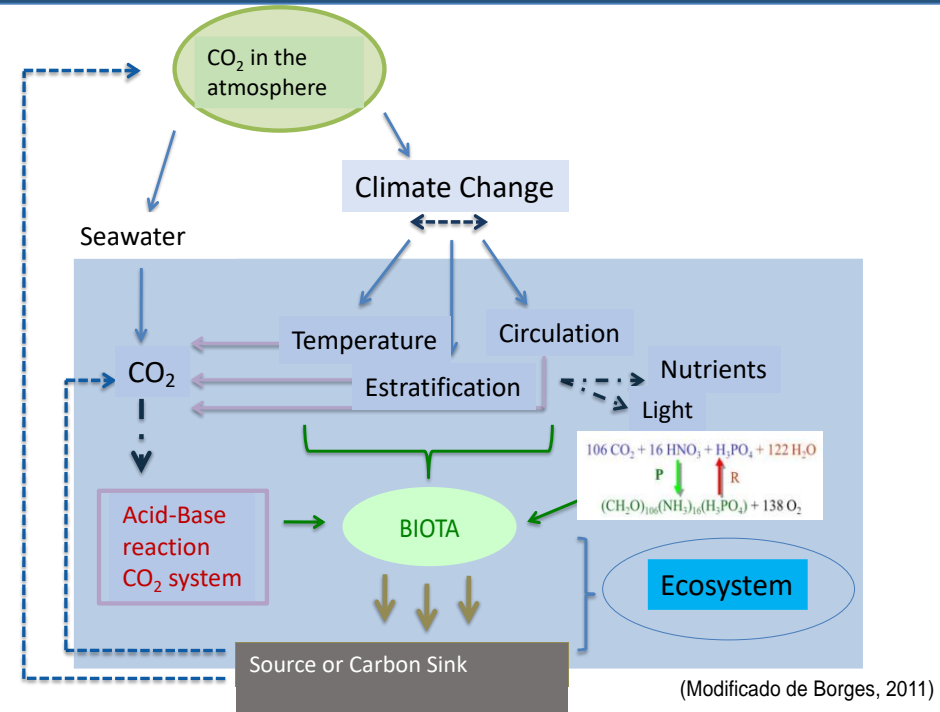
- Changes in pH and carbonate chemistry force marine organisms to spend more energy forming shells etc.
- Current rate of these chemistry changes are unprecedented (nothing like this in 300 million years)
- OA is **one more stressor** (cumulates with higher temperatures and deoxygenation)

OA BASICS

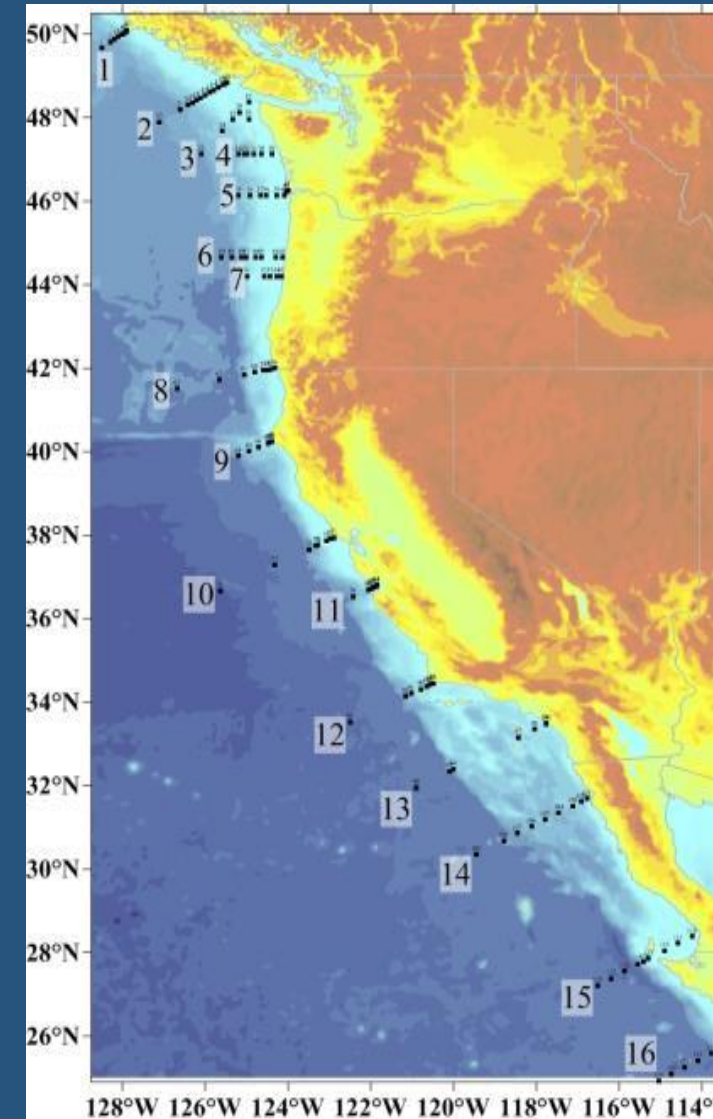
The Carbon Cycle



The ocean offers us ecosystem services



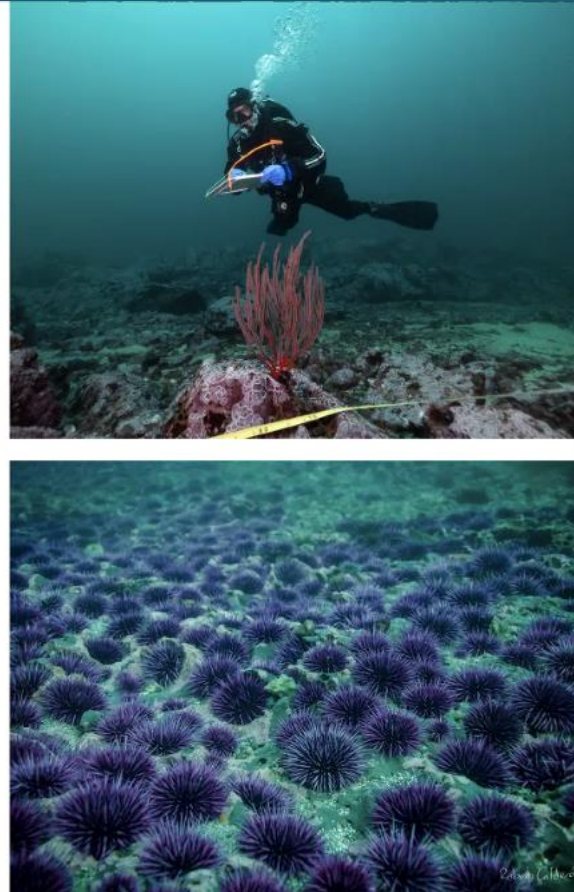
In 2007 Coastal of Baja California Mexico use to see like this photo





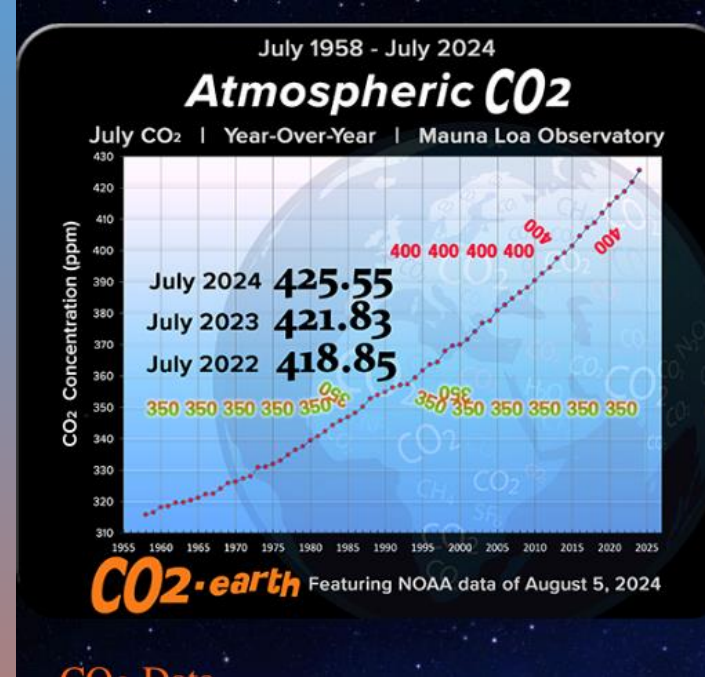
Dra. Fiorenza Michelli

The same environment after 2014

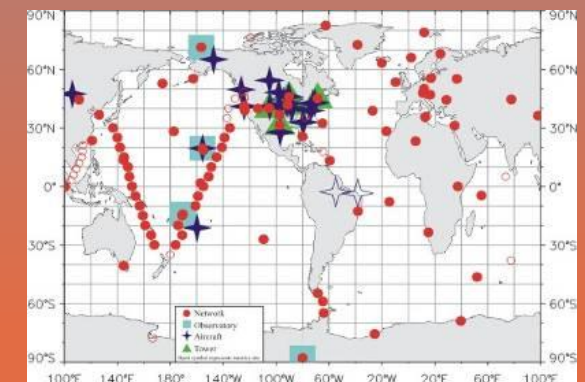
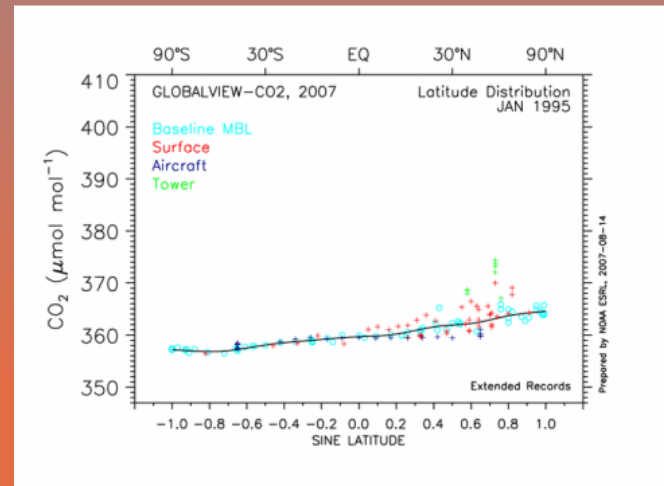


Marine forests, between collapse and reforestation plans
Giant algae create marine forests on the northeastern Pacific coasts, but in Baja California these ecosystems are threatened due to heat waves and purple urchin infestations.

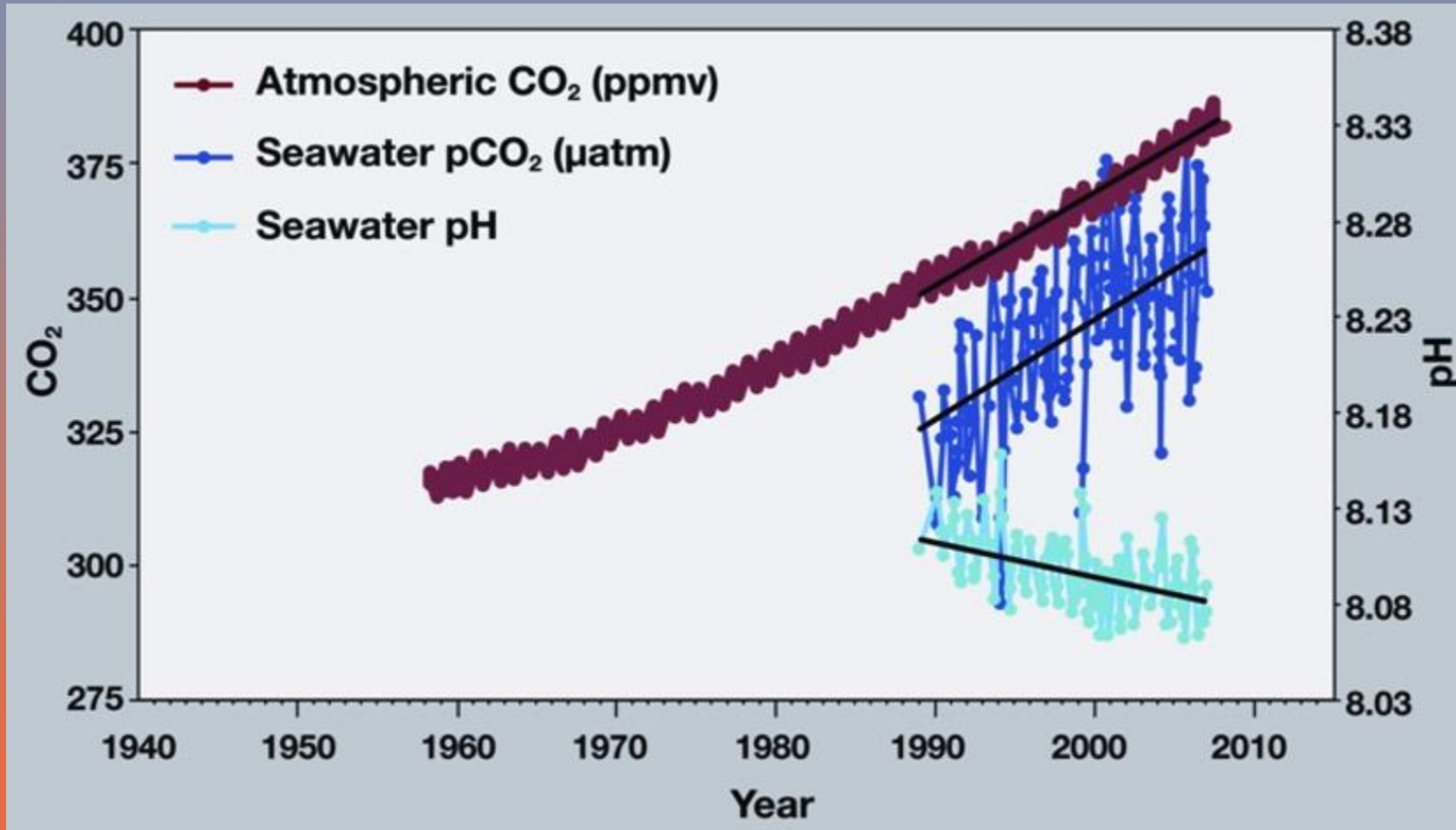
- Atmospheric CO₂ has risen by about 40% above pre-industrial levels.
- The ocean absorbs about 1/3 of human-caused emissions of CO₂ annually, thereby making seawater more acidic (decreasing pH).



26%
 $8.8 \pm 1.8 \text{ GtCO}_2/\text{yr}$



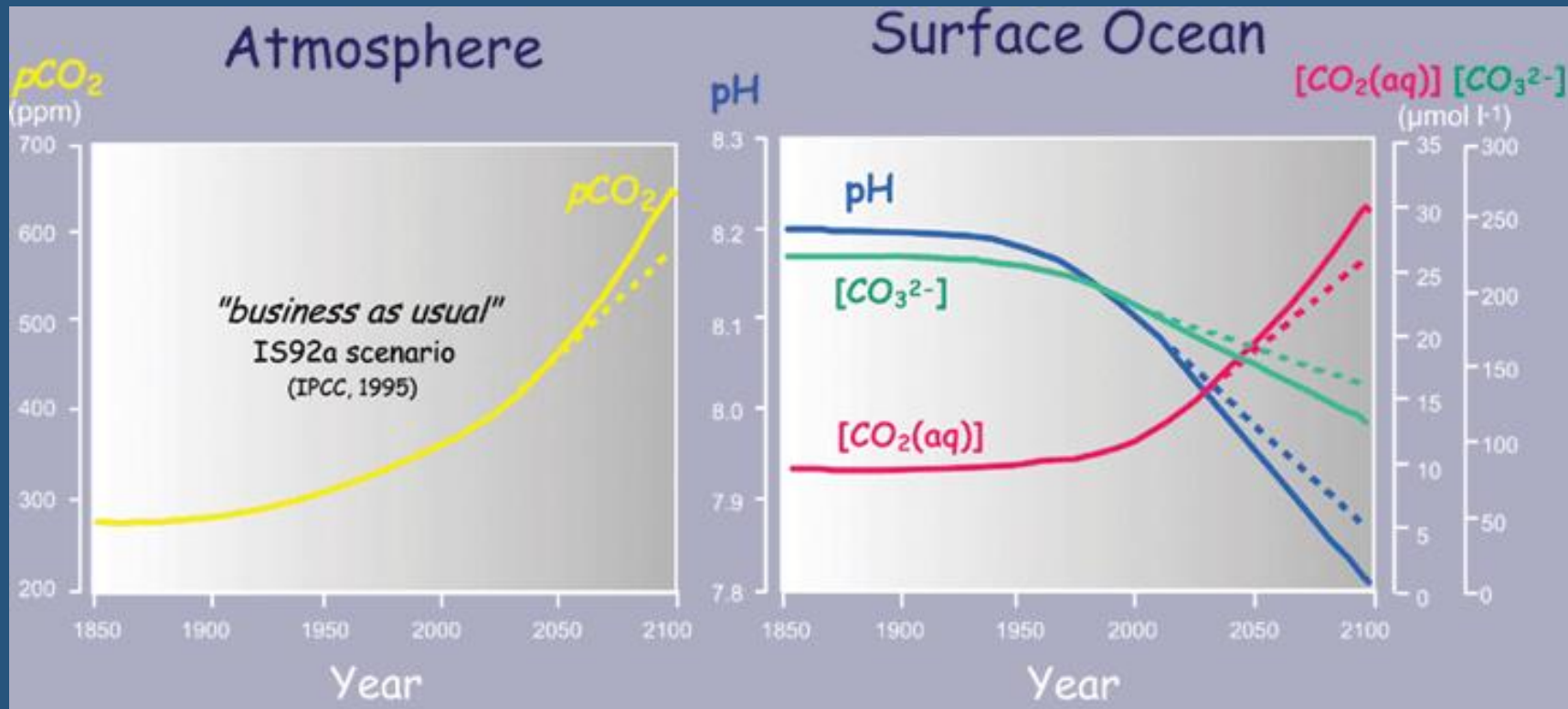
Change in Ocean pH mirrors CO₂ Concentrations (pCO₂)



pCO₂ (CO₂ partial pressure) is commonly referred to as CO₂ concentration in seawater

Image courtesy of NOAA. Modified after R. A. Feely, Bulletin of the American Meteorological Society, July 2008.

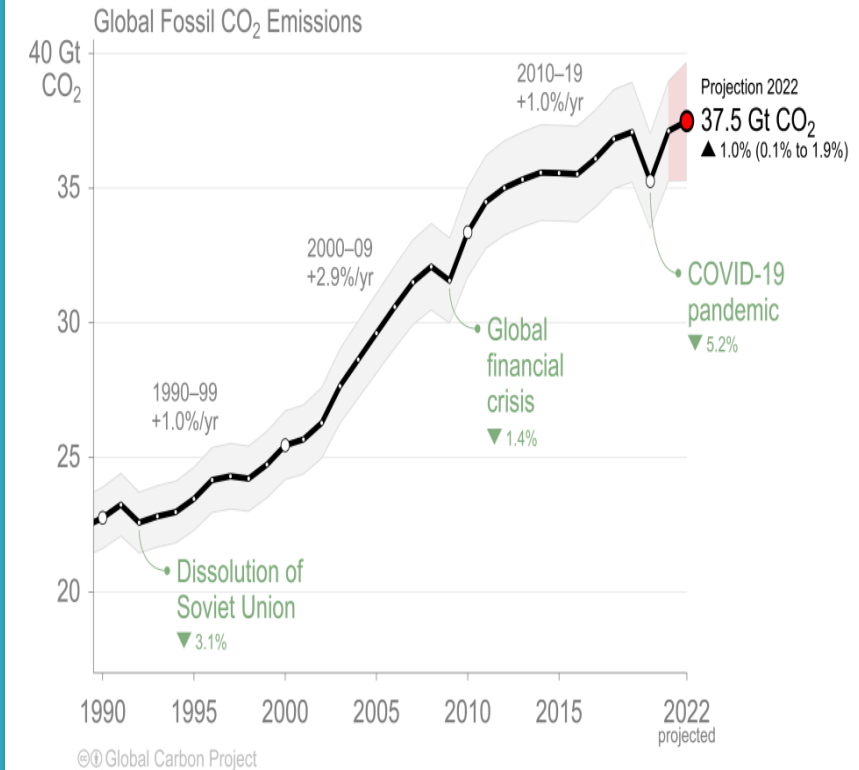
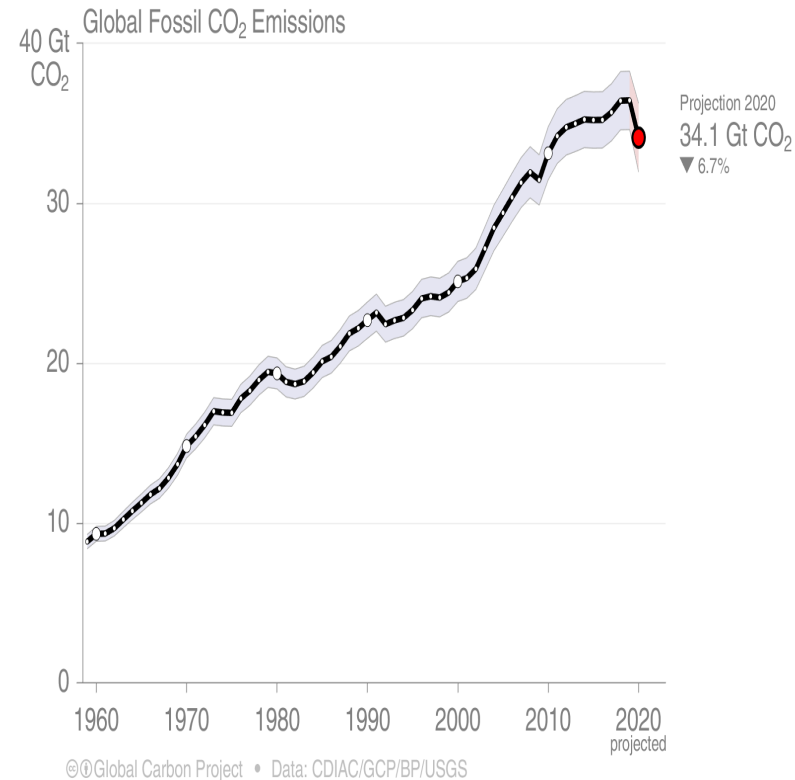
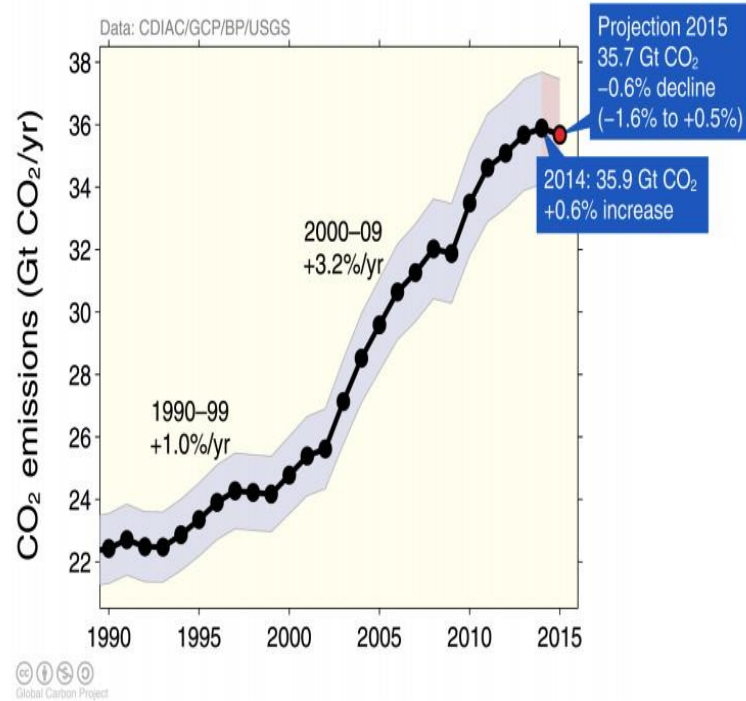
Projections show that continued emissions could drop ocean pH another 0.5 pH units to 7.7 by 2100. A decrease of 60% for CaCO_3 concentration is expected.



Tomado de Wolf-gladrow et al., 1999.

Fossil Fuel Emissions: Actual vs. IPCC Scenarios

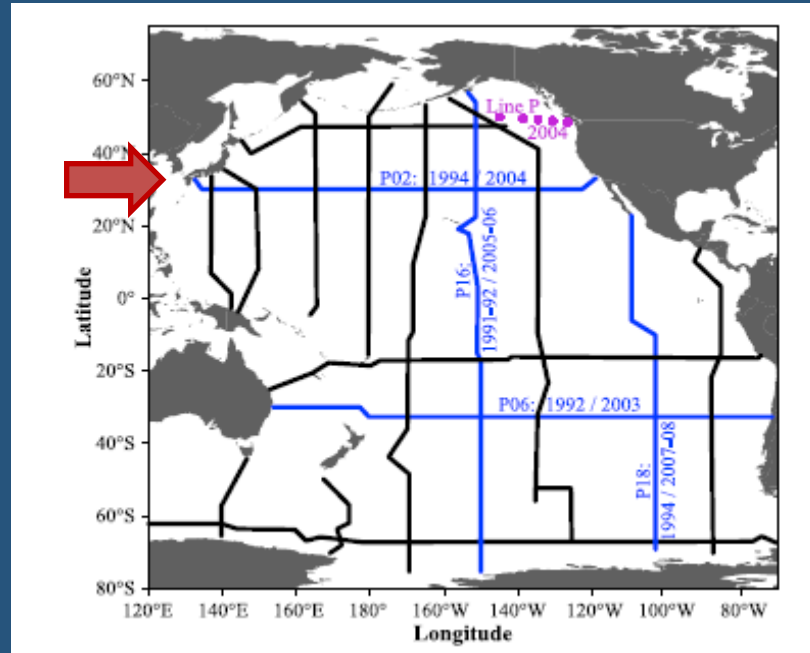
Emissions from fossil fuel use and industry



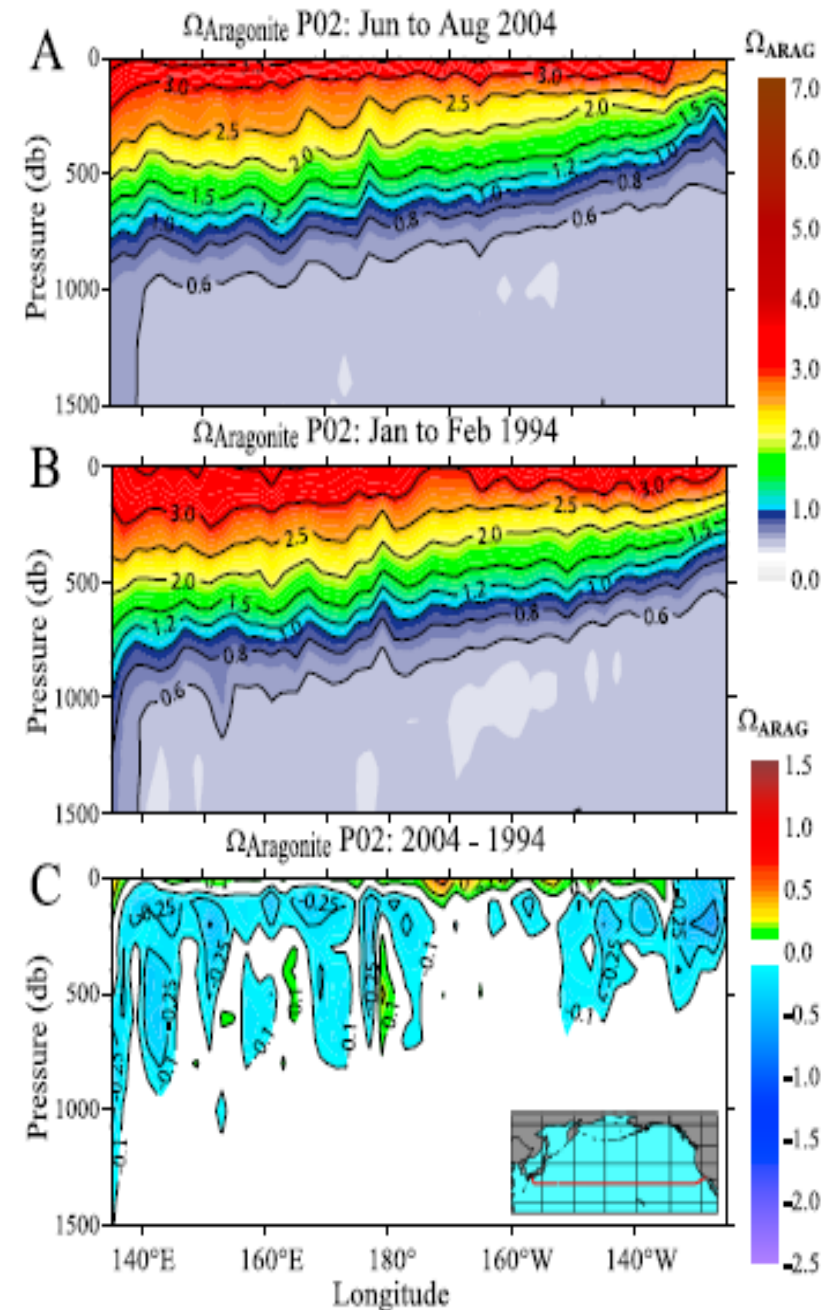
Decadal changes in the aragonite and calcite saturation state of the Pacific Ocean

Richard A. Feely,¹ Christopher L. Sabine,¹ Robert H. Byrne,² Frank J. Millero,³ Andrew G. Dickson,⁴ Rik Wanninkhof,⁵ Akihiko Murata,⁶ Lisa A. Miller,⁷ and Dana Greeley¹

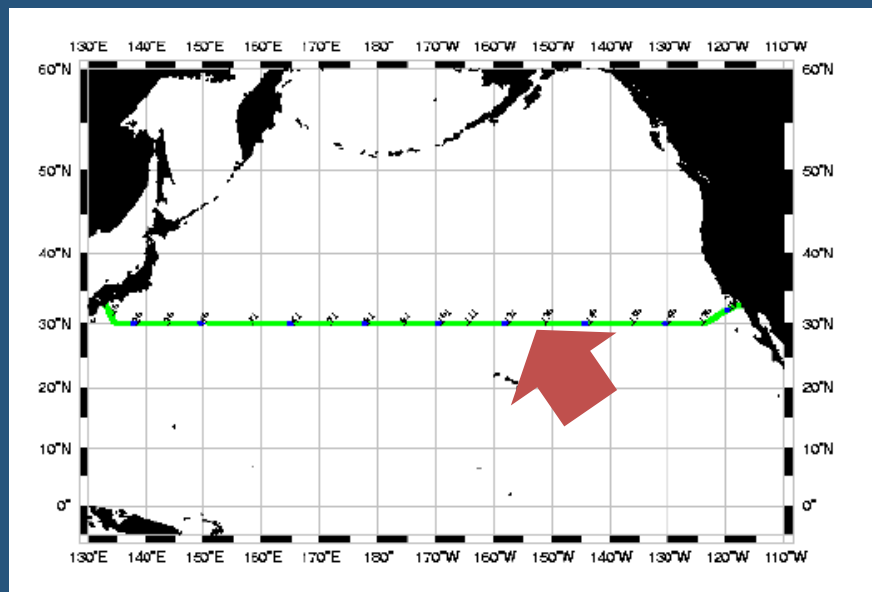
GLOBAL BIOGEOCHEMICAL CYCLES



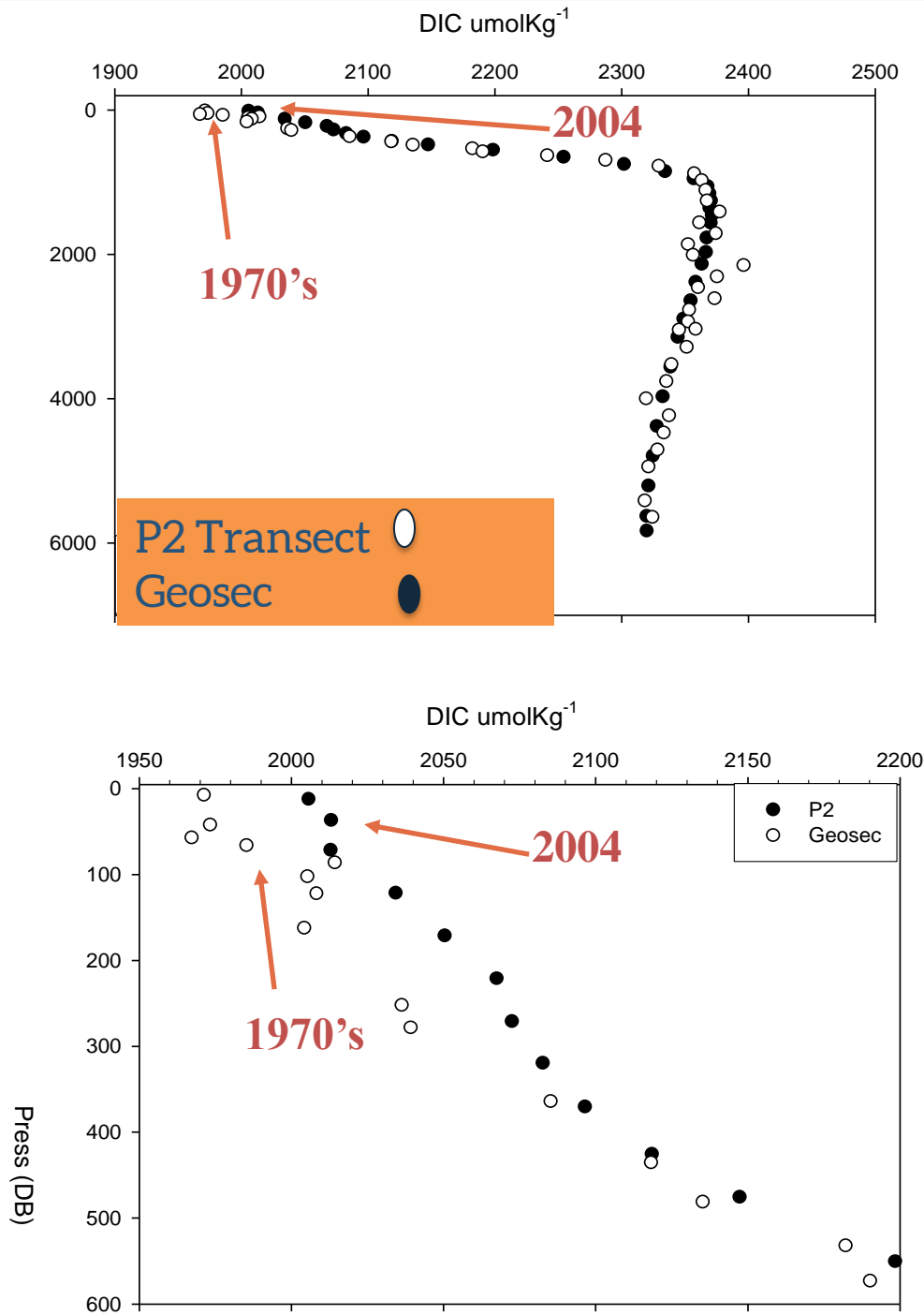
The saturation horizon
It's getting shallower
From 1 to 2m per year.



“See to be believe”



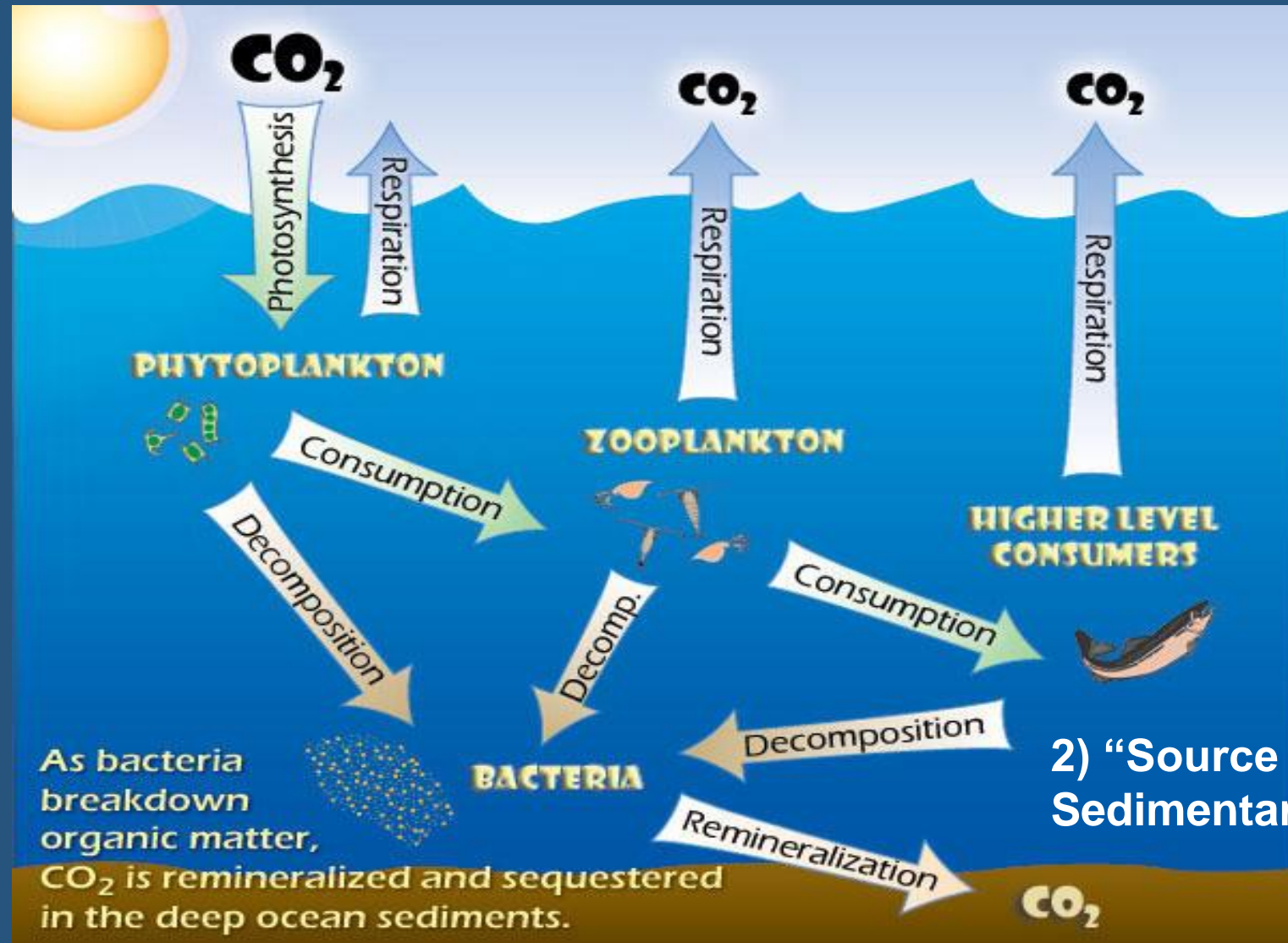
R/V Roger Revelle
is operated by
Scripps Institution
of Oceanography



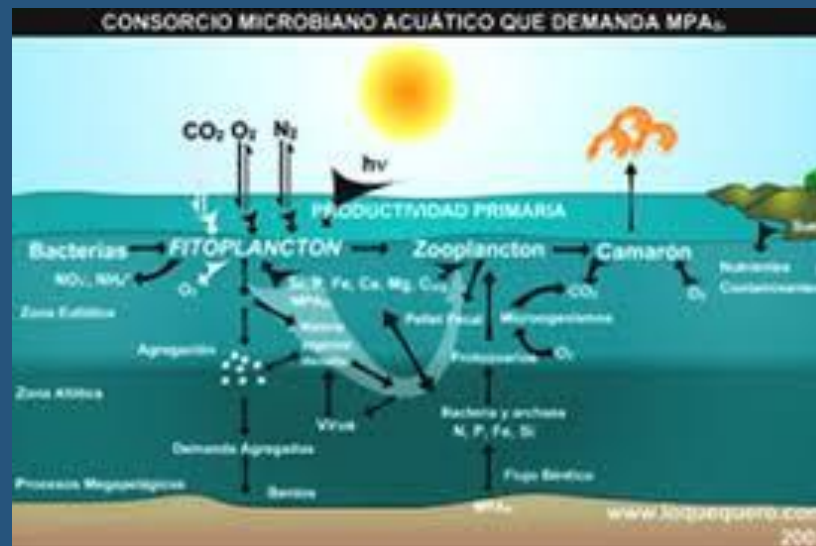
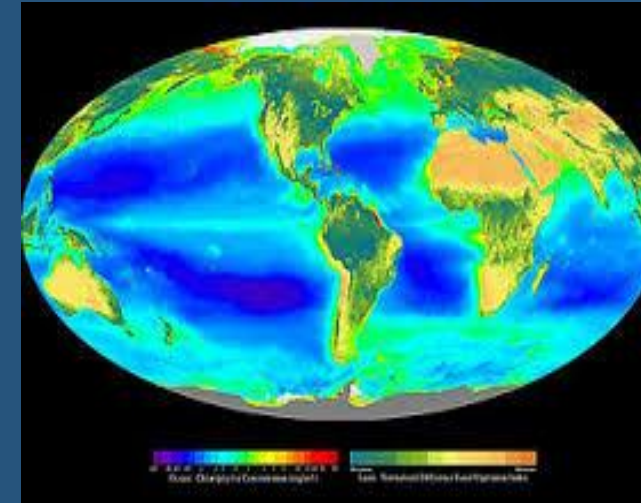
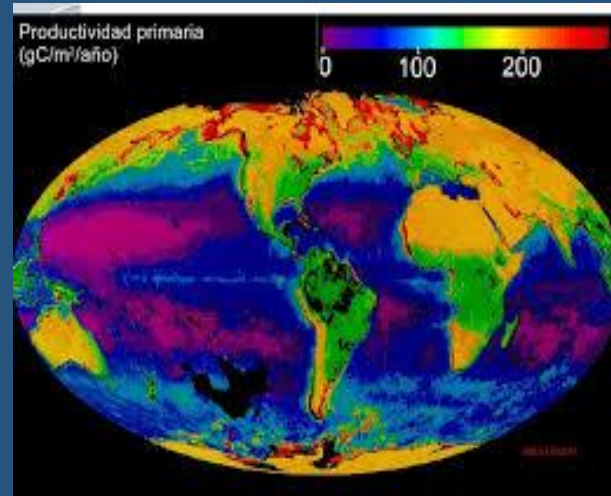
“¿Why is important understand the CO₂ System ?

- CO₂ controls the fraction of incoming radiation that remains trapped in the atmosphere (greenhouse effect), which controls the planetary climate.
- CO₂ is the raw material used to build organic matter.
- CO₂ controls the pH of the oceans
- The distribution of CO₂ species affects the preservation of CaCO₃ deposited on the seabed

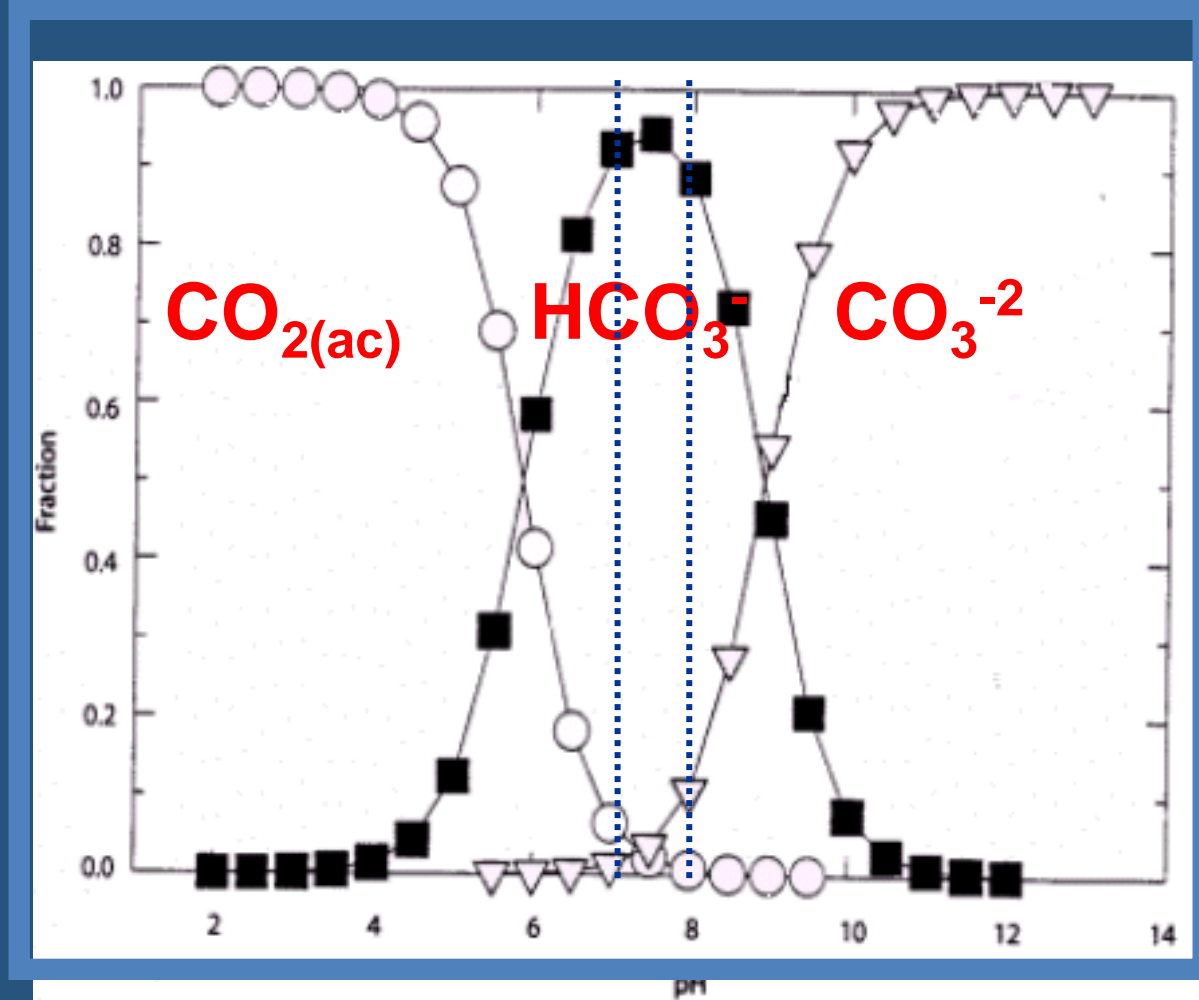
Controls the amount of CO₂ in the atmosphere



Carbon Source for Primary Organic Productivity

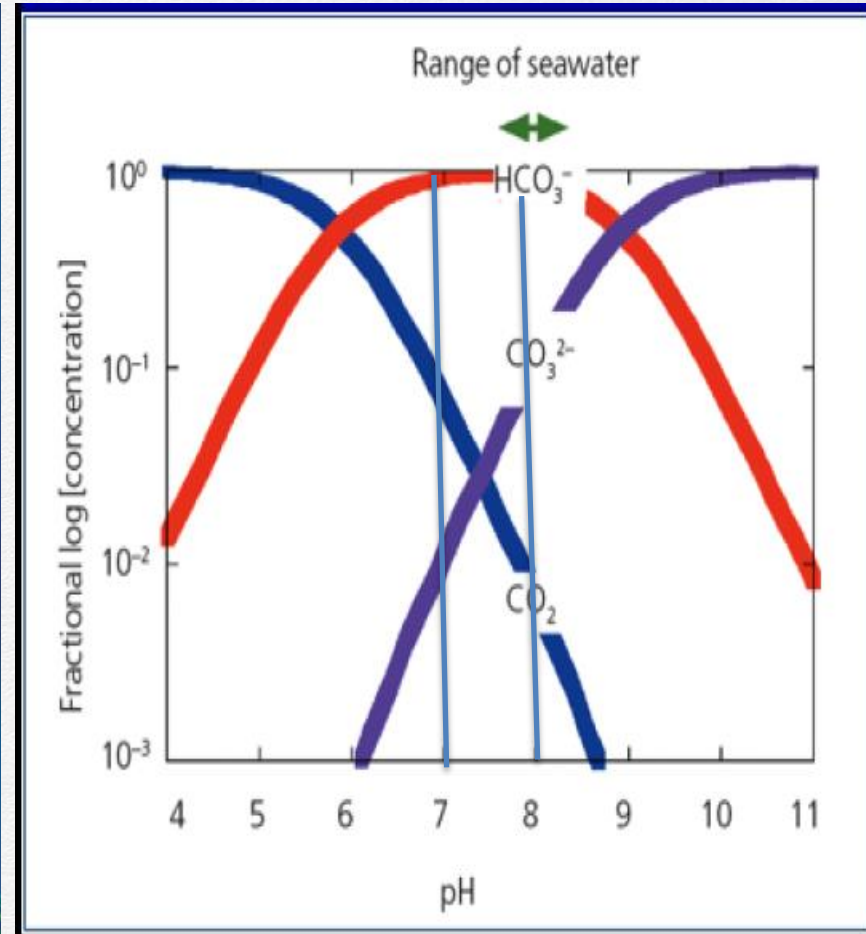
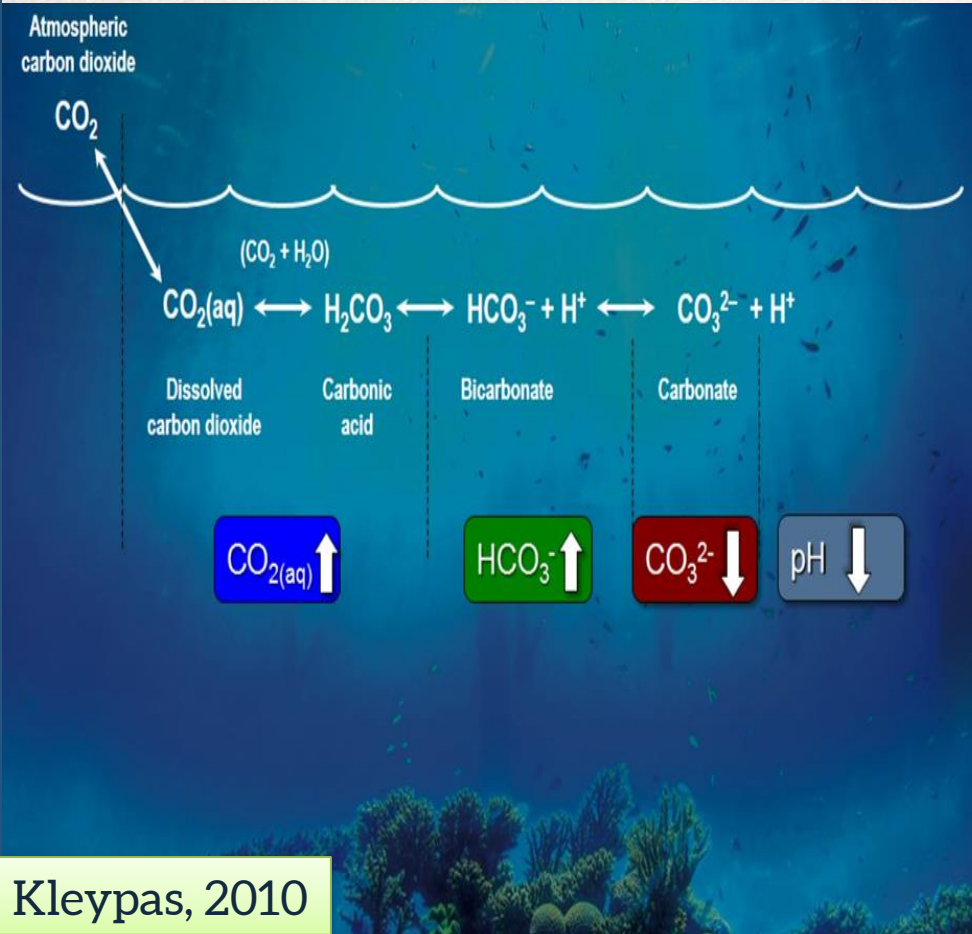


“Regulates pH”



Tomado de Millero et al., 1994

Therefore the term “ocean acidification” is the decrease in pH in seawater due to the dissolution of anthropogenic CO_2 .





To study the Ocean Acidification “OA” several “Tools” are required:

1. Understanding acid-base reactions in seawater
“Especially CO_2 ”
 2. Understand which CO_2 parameters (pH, Total Alkalinity, Dissolved Inorganic Carbon, pCO_2) are normally measured in seawater and how.
 3. Have access to a computing tool to perform CO_2 system calculations such as CO2Sys or CO_2 Cal.
 4. Have access to analytical equipment (and training) to perform measurements of carbon variables.
-



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-

Part 1: Seawater carbonate chemistry

Available in the web

1 The carbon dioxide system in seawater: equilibrium chemistry and measurements

Andrew G. Dickson

Scripps Institution of Oceanography, University of California, USA

1.1 Introduction

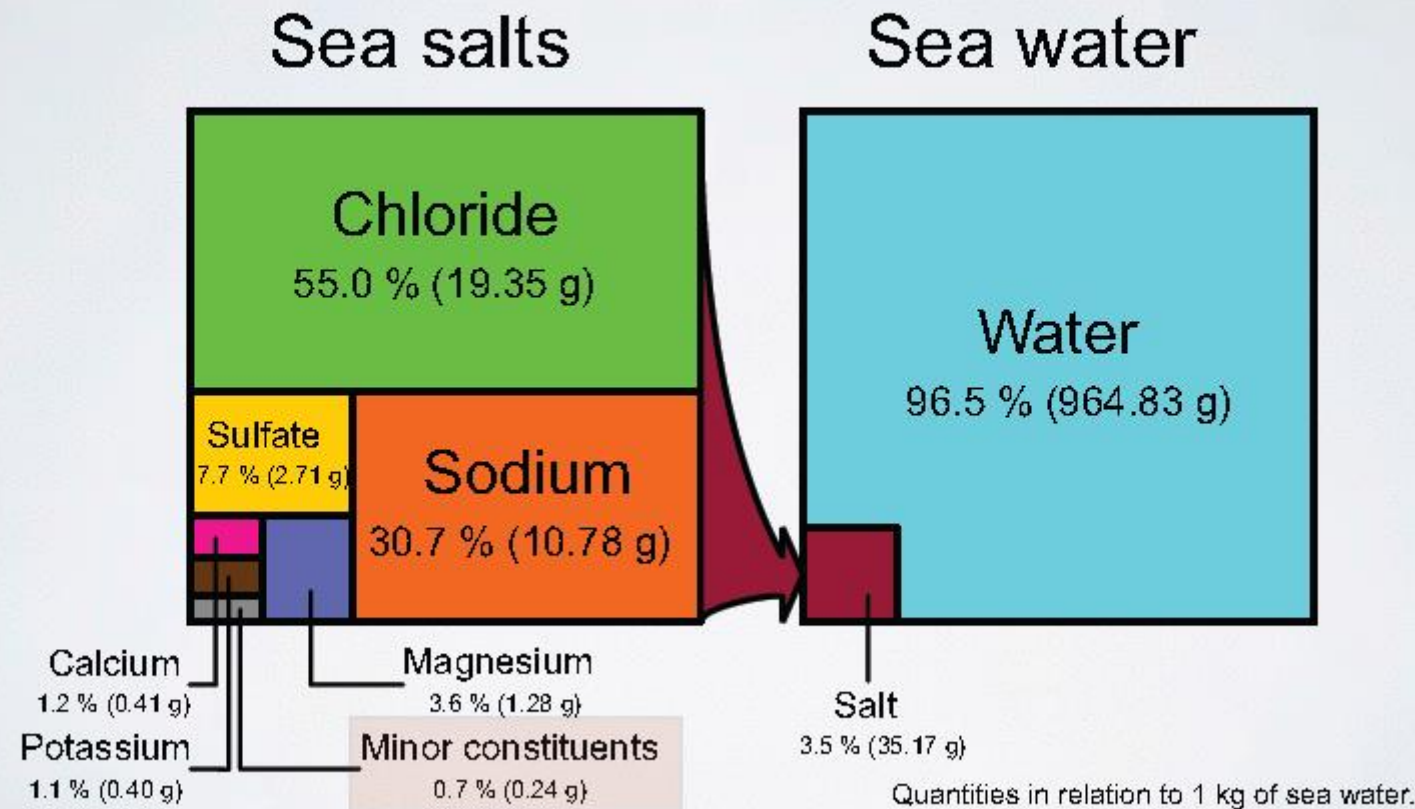
The world's oceans can be thought of as a dilute solution of sodium bicarbonate (together with species at still lower concentrations) in a saltwater background. In the surface water, for example, the concentration of total dissolved inorganic carbon (the sum of the coexisting species: bicarbonate ion, carbonate ion, and unionised dissolved carbon dioxide) is about 2 mmol kg^{-1} . About 90% of this is present as bicarbonate ion, the proportion of carbonate ion is about 10% less ($\sim 10\%$), and that of unionised carbon dioxide yet another factor of 10 less. The equilibria between these various species (see below), seawater is buffered (weakly) with respect to hydrogen ion (present at much lower concentrations: $<10^{-8} \text{ } \mu\text{mol kg}^{-1}$).

Over the past twenty years, accurate measurement of the seawater carbon dioxide system has become a high priority for scientists who have worked to understand just how much of the carbon dioxide from human activities has ended up in the ocean, where it is distributed, and how it has changed the chemistry of the oceans. The chemical changes associated with the increase of CO_2 in the oceans are known as *ocean acidification*. As we work to design suitable experiments to understand the basic consequences of such changes, it is important that the chemistry of CO_2 be well characterised from both laboratory experiments and field observations that are undertaken. Achieving this requires a good understanding of the basic solution chemistry underlying ocean acidification, as well as of the relative



INTRODUCTION TO CO₂ EQUILIBRIA IN SEA WATER

COMPOSITION OF SEA WATER ($S = 35$)

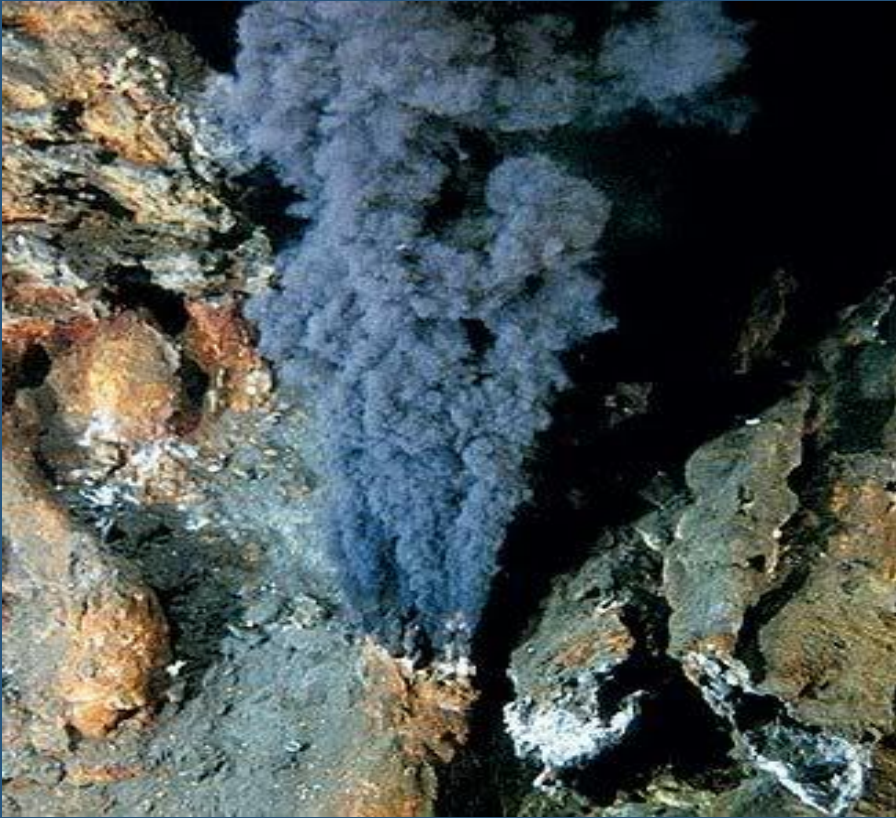


Why is salty?



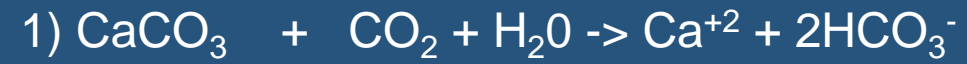
“Not salinity changes since 200 millions of years ”

Hydrothermalism and volcanism

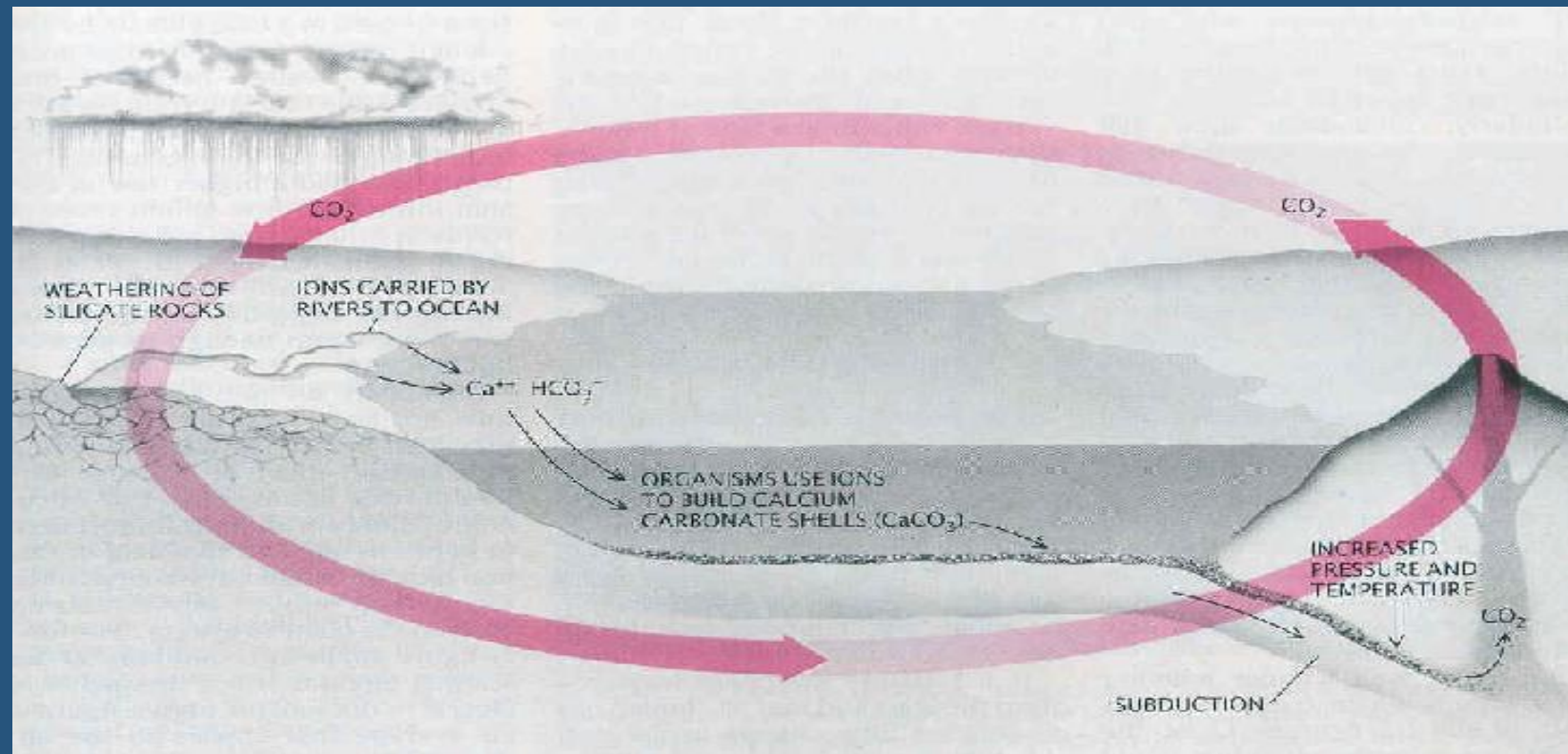



“Rivers?”

Rocks lweathering:

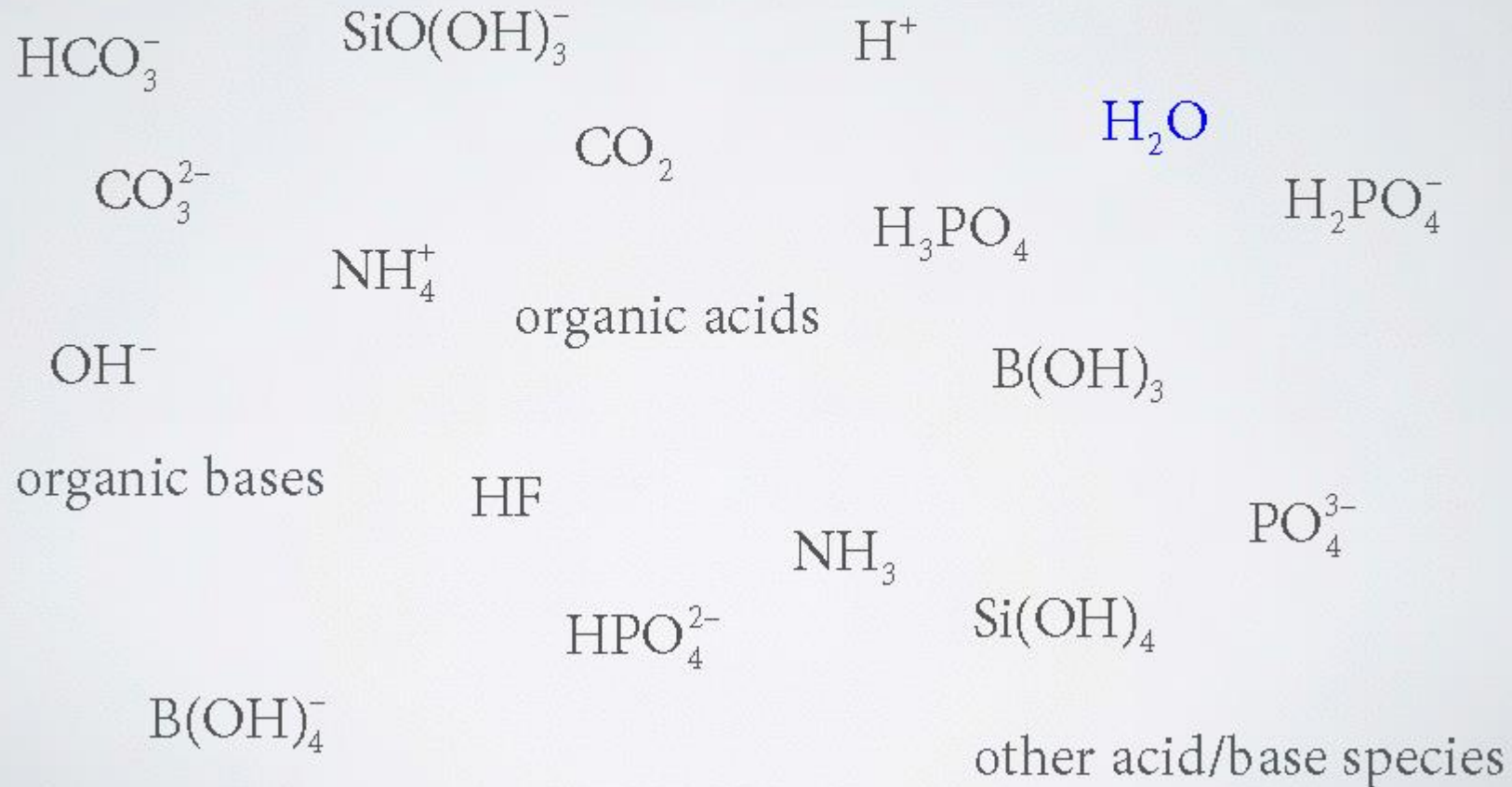


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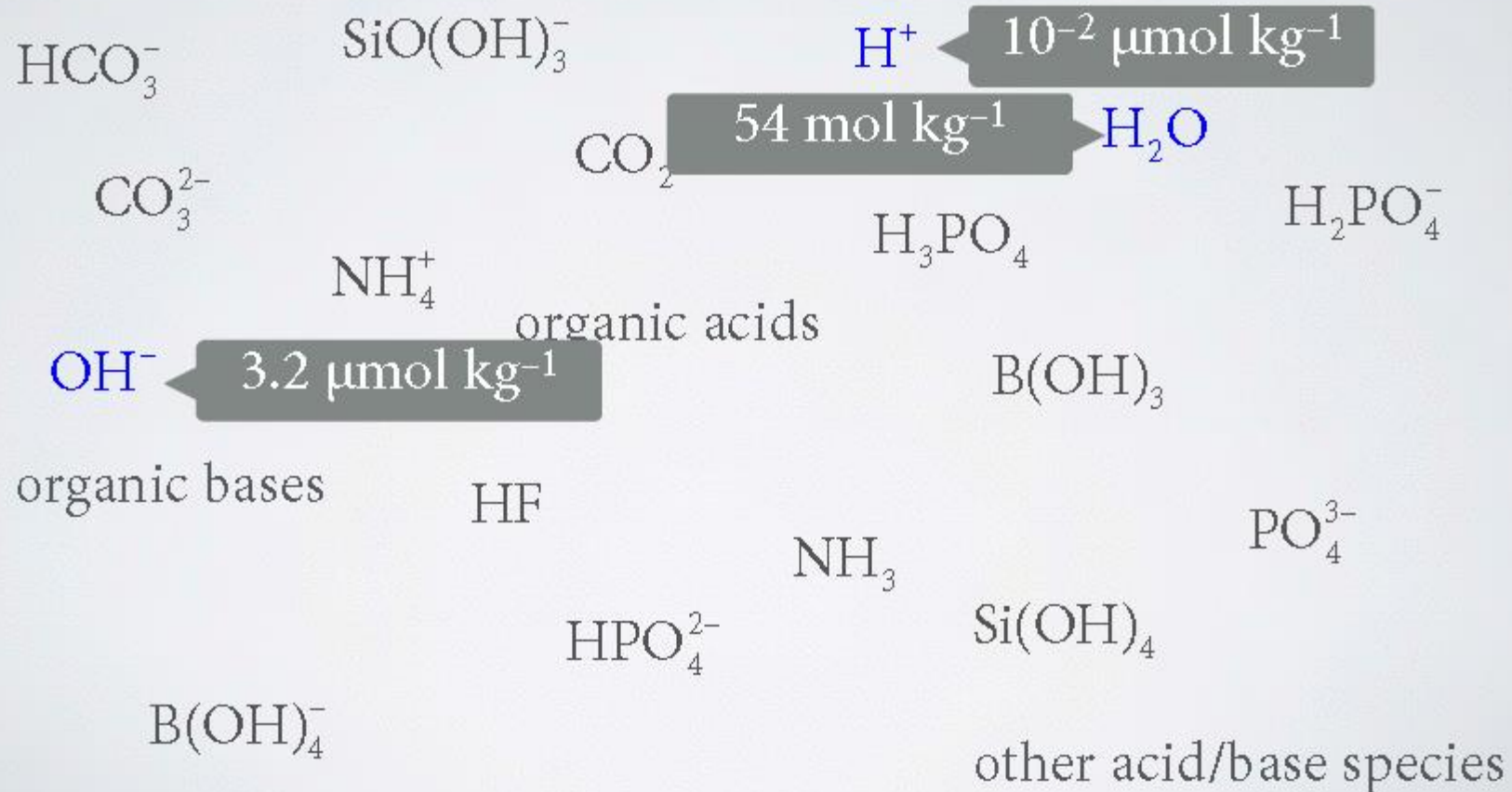


ACID-BASE SPECIES COMPRISE A SUBSET OF THE “MINOR CONSTITUENTS” PRESENT IN SEAWATER



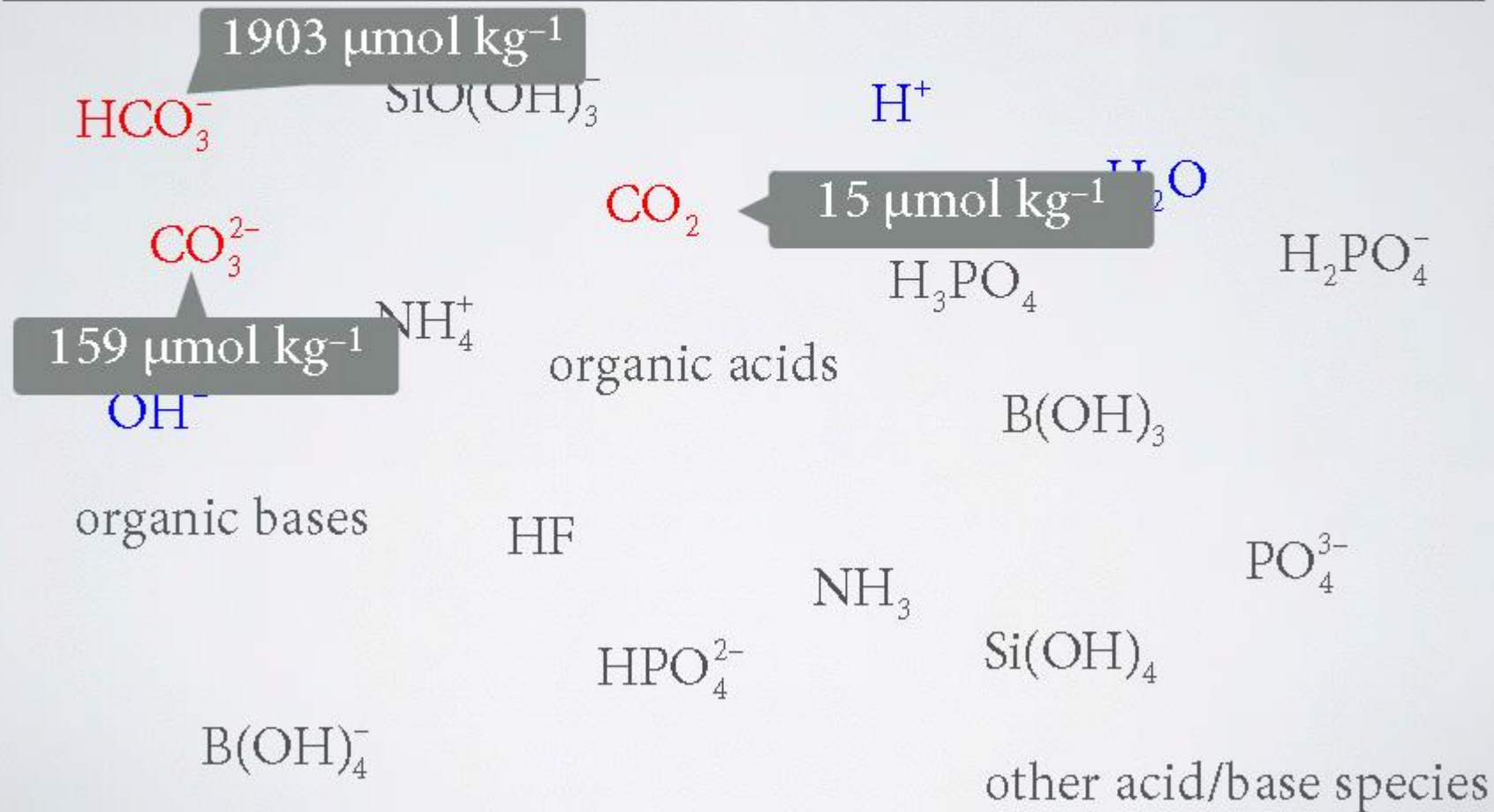
Average ocean water; pH = 8; $t = 18\text{ }^{\circ}\text{C}$

ACID-BASE SPECIES COMPRISE A SUBSET OF THE “MINOR CONSTITUENTS” PRESENT IN SEAWATER



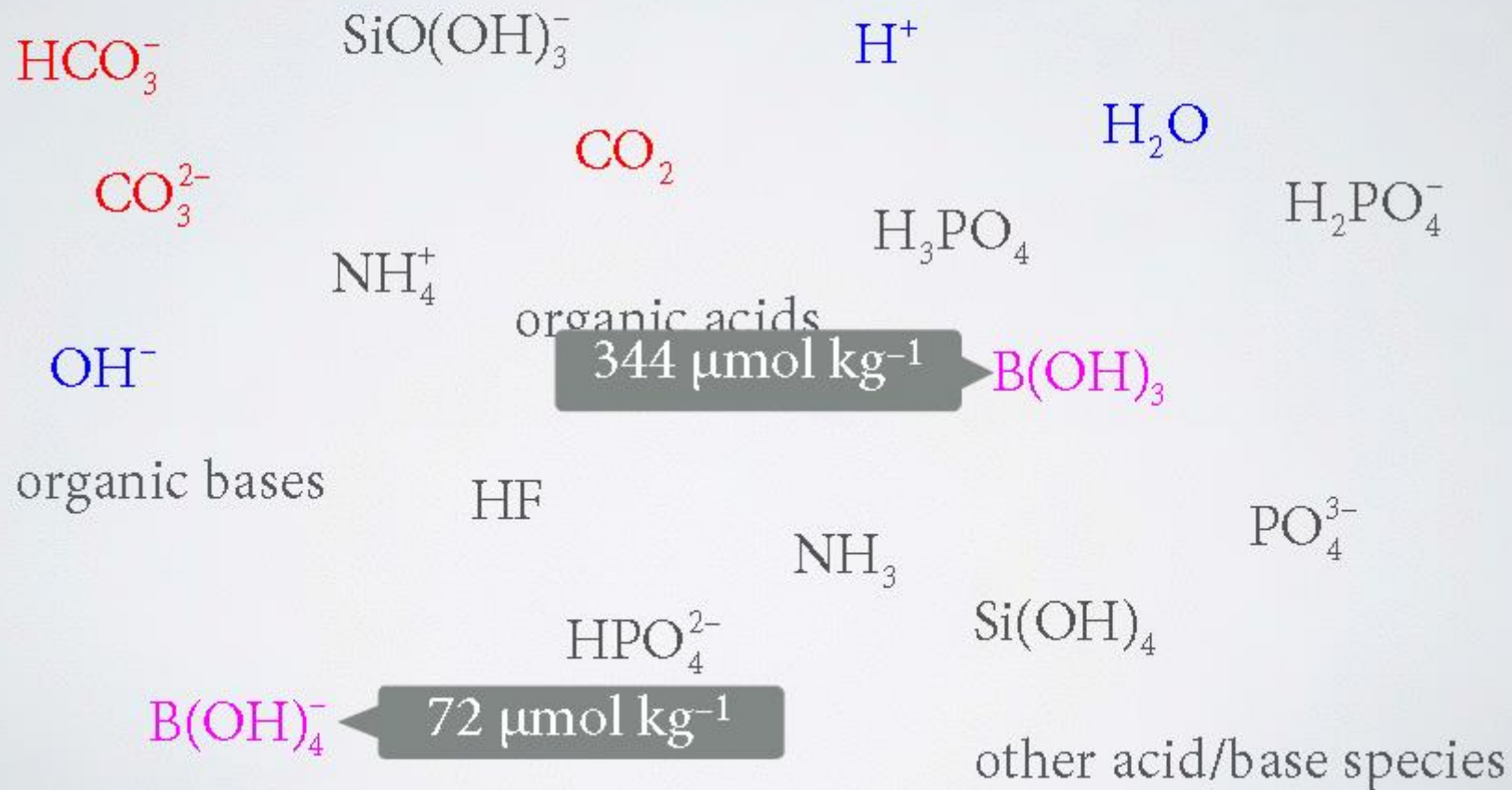
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ACID-BASE SPECIES COMPRISE A SUBSET OF THE
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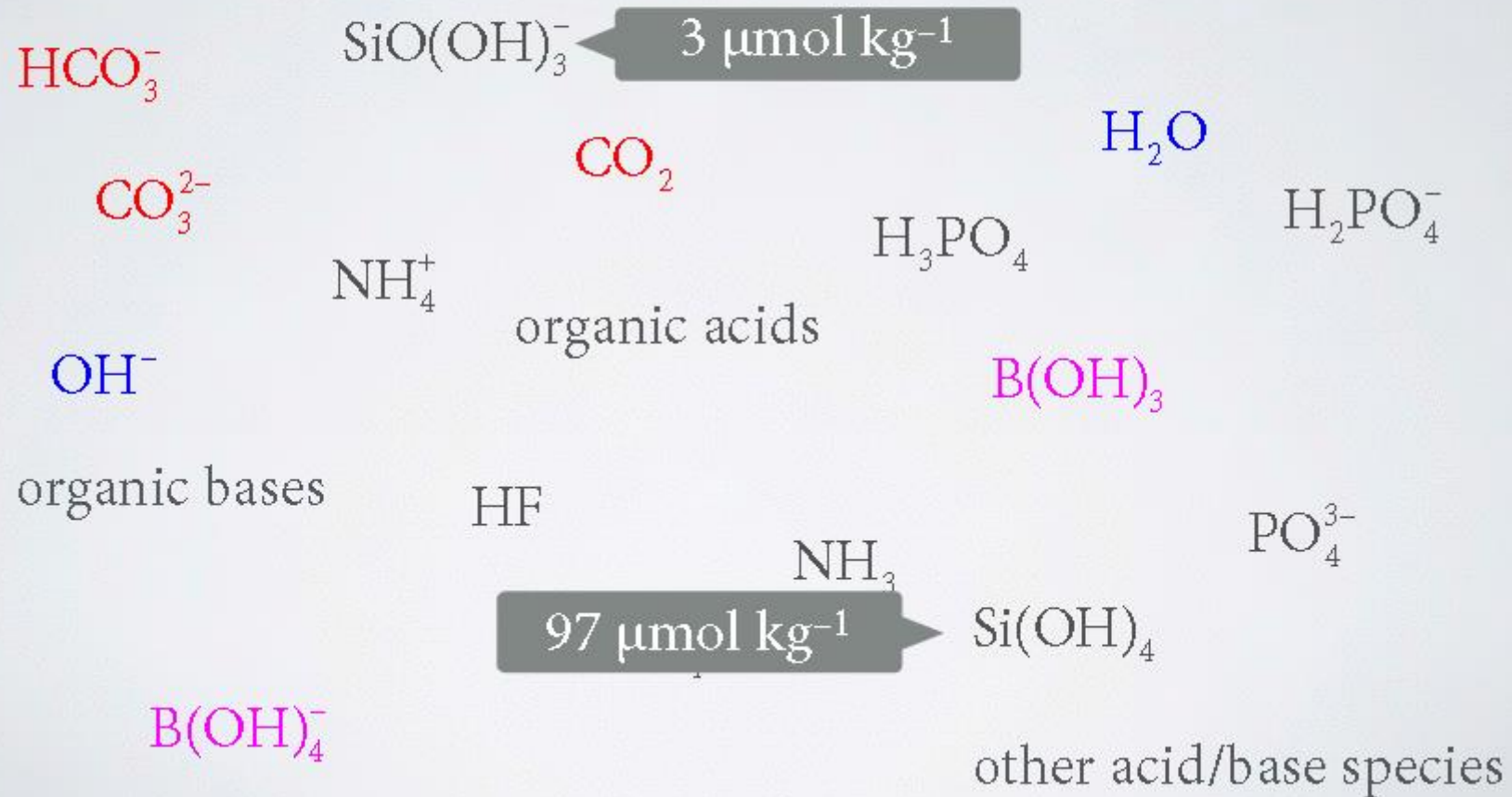
Average ocean water; pH = 8; $t = 18\text{ }^{\circ}\text{C}$

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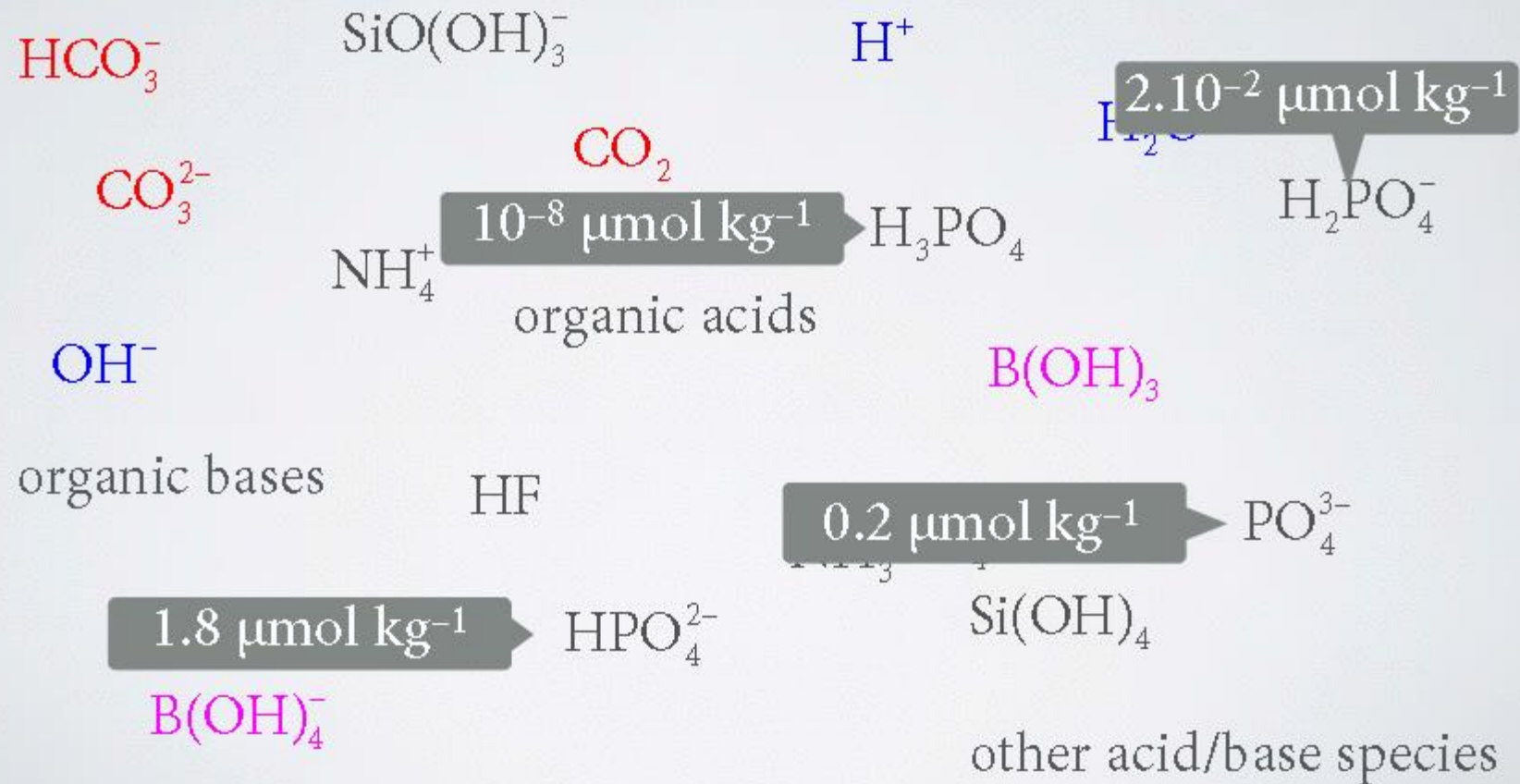
Average ocean water; pH = 8; $t = 18\text{ }^{\circ}\text{C}$

ACID-BASE SPECIES COMPRISE A SUBSET OF THE
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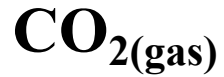


Average ocean water; pH = 8; $t = 18\text{ }^{\circ}\text{C}$

ACID-BASE SPECIES COMPRISE A SUBSET OF THE “MINOR CONSTITUENTS” PRESENT IN SEAWATER



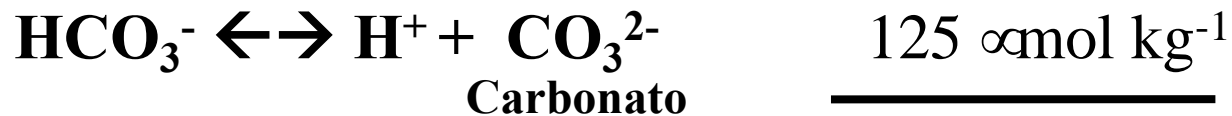
Carbon Chemistry in the Ocean and the challenge



"1 Atmosphere-Ocean Rx"



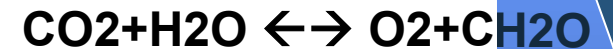
"2 Acido-Base Rx"



TCO₂ or DIC

2170 $\mu\text{mol kg}^{-1}$

"3 Oxido-Red"



"4 Precipitation-Dissolution Rx"



To study the Ocean Acidification “OA” several “Tools” are required:

1. Understanding acid-base reactions in seawater
“Especially CO_2 ”
 2. Understand which CO_2 parameters (pH, Total Alkalinity, Dissolved Inorganic Carbon, pCO_2) are normally measured in seawater and how.
 3. Have access to a computing tool to perform CO_2 system calculations such as CO2Sys or CO_2 Cal.
 4. Have access to analytical equipment (and training) to perform measurements of carbon variables.
-

We need at least two parameters of the CO₂ system to be able to estimate the other two and the different CO₂ species from the system

- Dissolved Inorganic Carbon
- pH
- pCO₂
- Total Alkalinity

temperature, salinity, pressure is needed

Options: silicate and phosphate.

Dissolved Inorganic Carbon

$$\text{DIC or TCO}_2 = [\text{H}_2\text{CO}_3] + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$



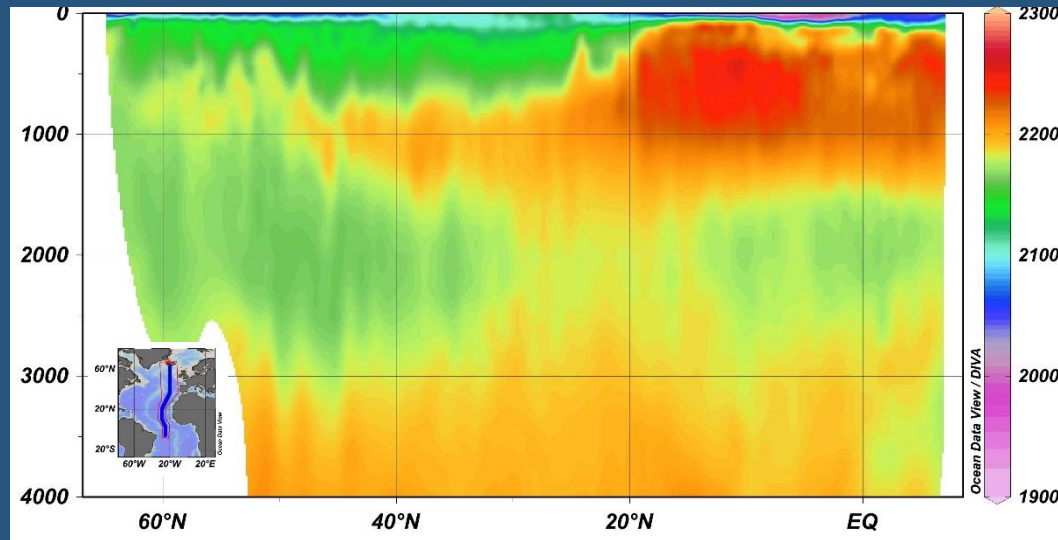
Sample (volume or mass)
+
 H_3PO_4

$\Rightarrow \text{CO}_2 \text{ (ac)}$

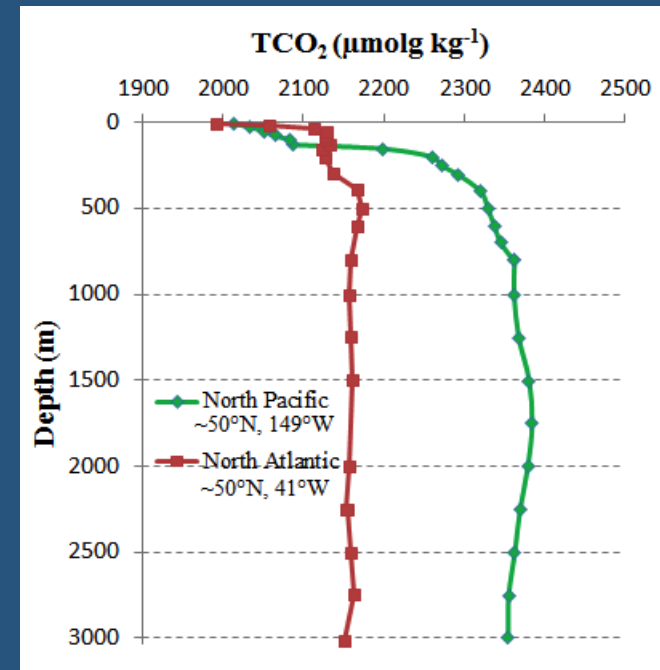
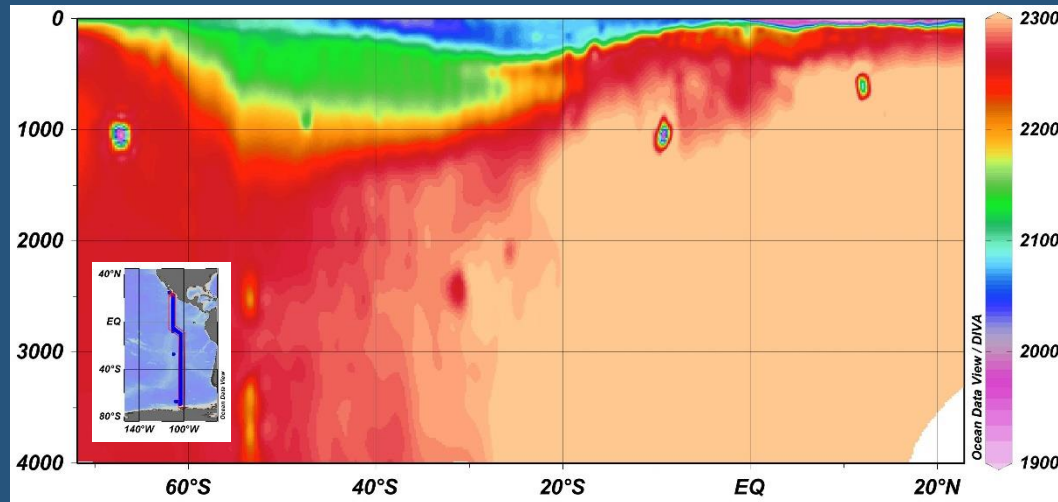
Determination by coulometry
(precision $1 \mu\text{mol/kg}$)

Total CO₂ or Dissolved Inorganic Carbon

Atlantic Ocean



Pacific Ocean



pH Sensors

- **Potentiometric**
- **Spectrophotometric**
Potentiometric determination (precision 0.005 pH units)

Spectrophotometric determination (precision 0.001 pH units)



pH: Scales (1) National Bureau Standards (NBS)
(2) Total Protons (H_{tot})
(3) Seawater (SWS)

$$\text{pH} = -\text{Log} [H^+]$$

pH scales related by the follow equations:

$$-\text{pH}_{\text{NBS}} \quad aH = 10$$

The reference state for the NBS/IUPAC scale is the infinitely dilute solution, which is very useful in very dilute natural waters, such as rivers and lakes. However, this scale is not recommended for seawater, due to its high ionic strength (Dickson, 1984; Millero et al., 1993). Preferable α scale based on proton concentration.

pH: (2) Total Protons (H_{tot})

Total pH scale, pH_T , 1973 (modified by Dickson 1990)

It takes into account

- 1) the free protons +
- 2) contribution from the association of the SO_4^{2-} ion with H^+ to form HSO_4^-

$$\text{pH}_T = -\log[\text{H}^+]_T$$



$$K'_S = [\text{H}^+]_F [\text{SO}_4^{2-}] / [\text{HSO}_4^-]$$

$$[\text{H}^+]_T = [\text{H}^+]_F (1 + [\text{SO}_4^{2-}] / K'_S)$$

pH: (3) Seawater (SWS)

Seawater pH scale, pH_{SWS}, 1979

It takes into account

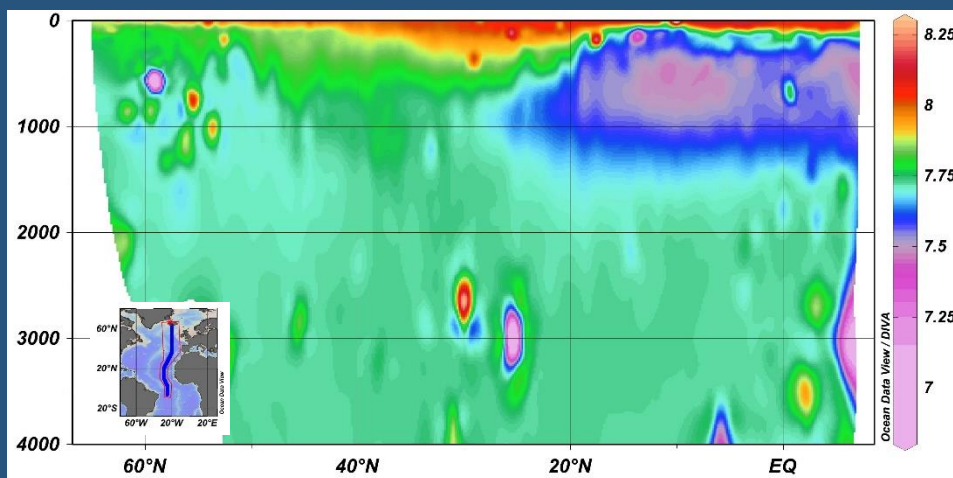
- 1) the free protons +
- 2) contribution from the association of the SO_4^{2-} ion with H^+ to form HSO_4^-
- 3) In addition to sulfates, it includes contributions from the fluoride ion.

$$\text{pH}_{\text{SWS}} = -\log[\text{H}^+]_{\text{SWS}} = -\log([\text{H}^+]_{\text{F}}(1 + [\text{SO}_4^{2-}]/K'_\text{S} + [\text{F}^-]/K'_\text{F}))$$

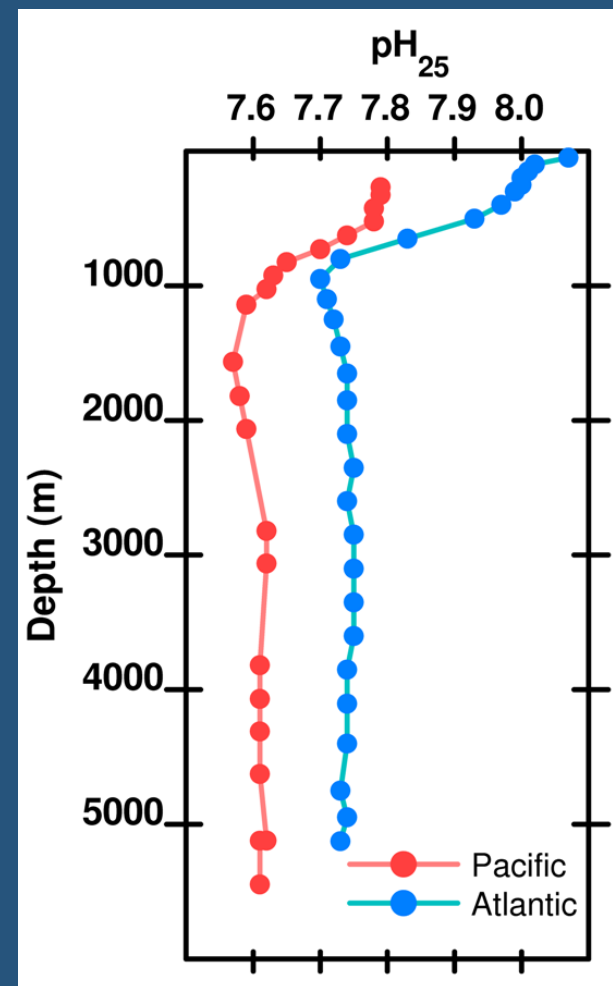
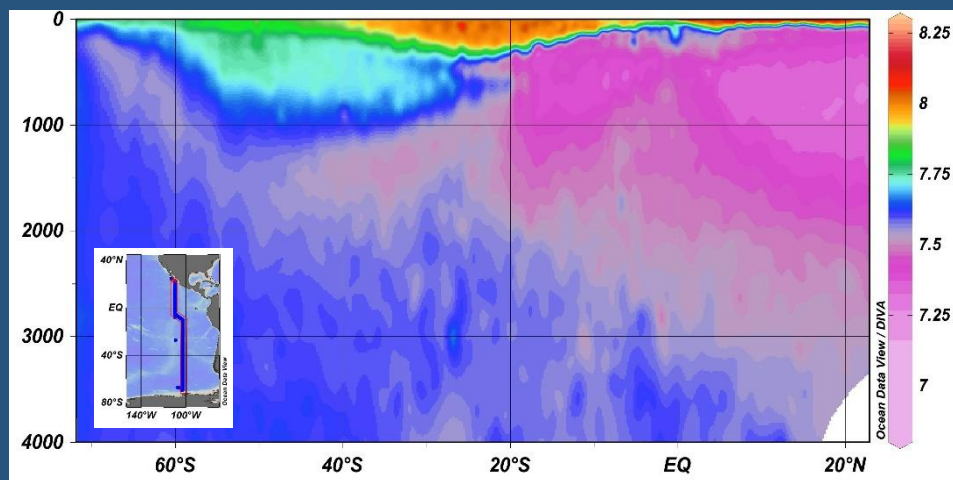
As a comparison important!!

To 20 deg C, Salinity 35, and 1 atm, pH values in protons total escale are aproximates:
.01 units higher in seawater scale
.13 units lower in NBS pH escale

Atlantic Ocean



Pacific Ocean



Partial Pressure of CO₂

¿What is pCO₂?

Pressure exerted by CO₂ on an inert gas (e.g. N₂) in equilibrium with a solution.

$$p\text{CO}_2 = x\text{CO}_2 (P_{\text{air}} - p\text{H}_2\text{O})$$

$x\text{CO}_2$: molar fraction (ppm, in dry air)

$p\text{H}_2\text{O}$: water vapor pressure

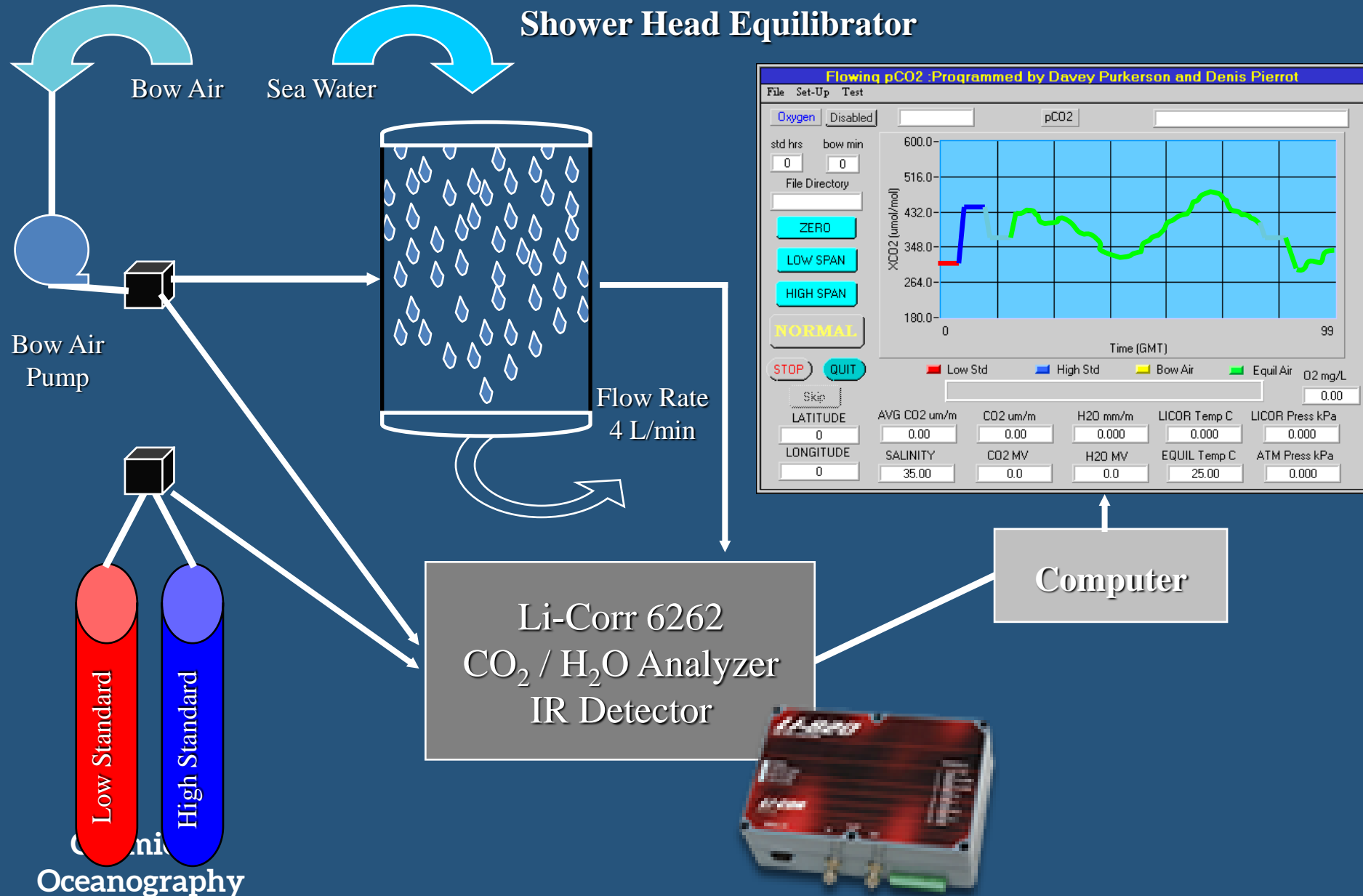
$$p\text{CO}_2 \approx f\text{CO}_2$$

Determination by infrared absorption spectroscopy (IR sensor)
(precision 1 μatm)

Spectrophotometric determination (precision 2-5 μatm)

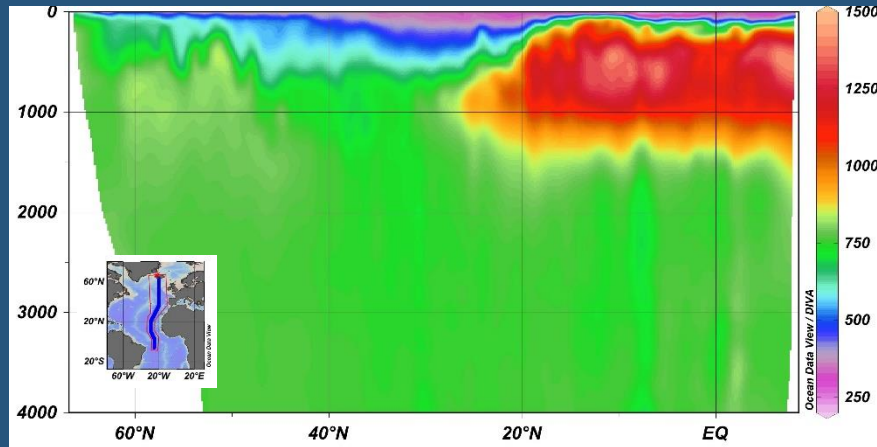


Flowing pCO₂ System

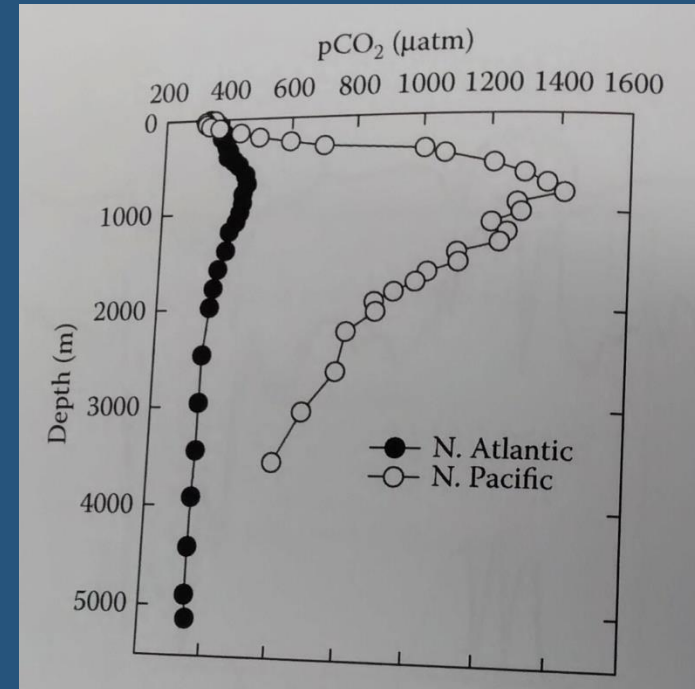
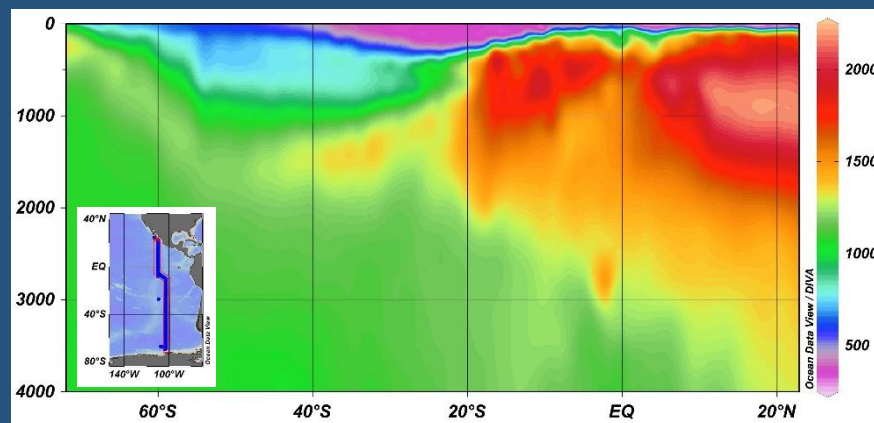


Parámetros del sistema del carbono: $p\text{CO}_2$

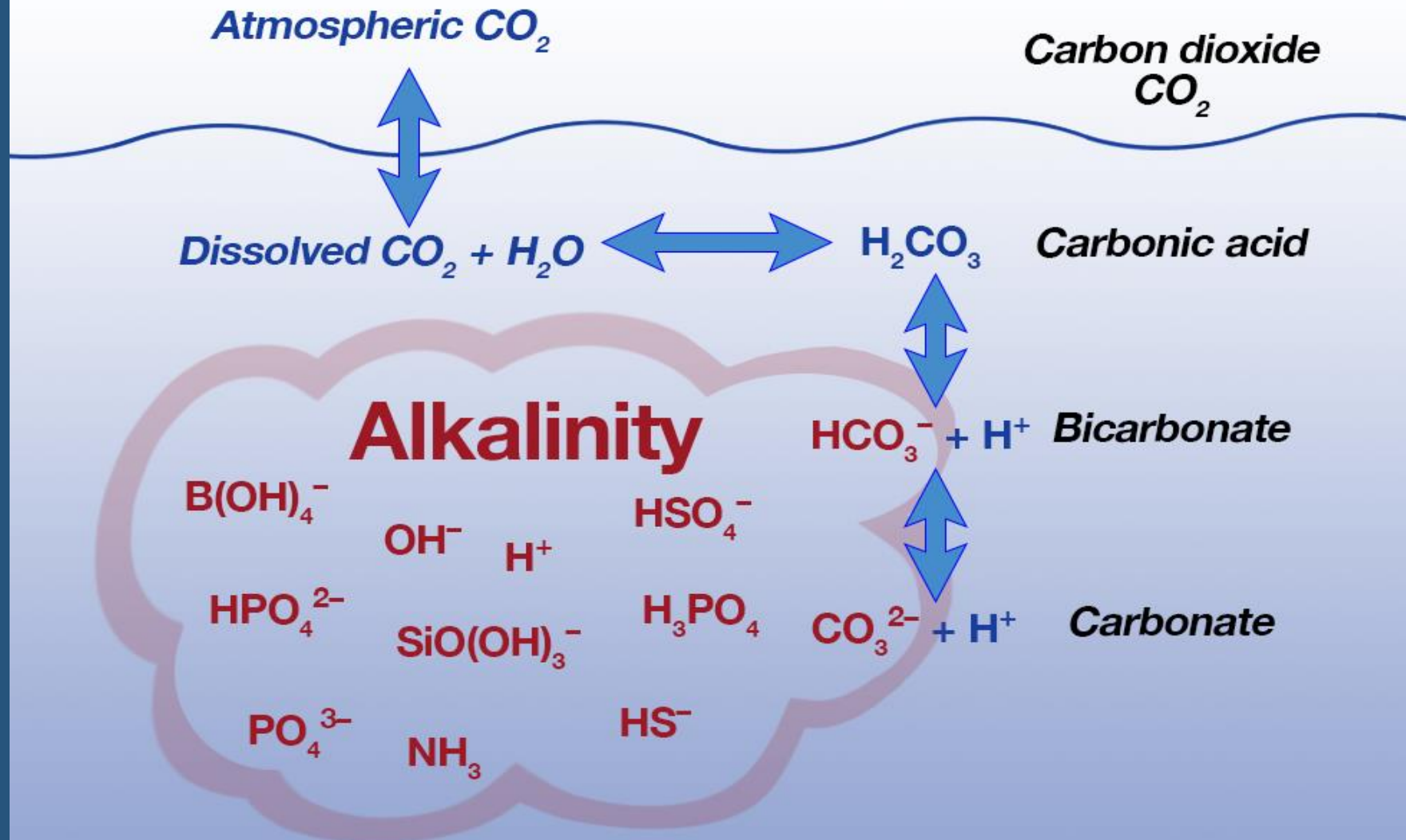
Atlántic Ocean



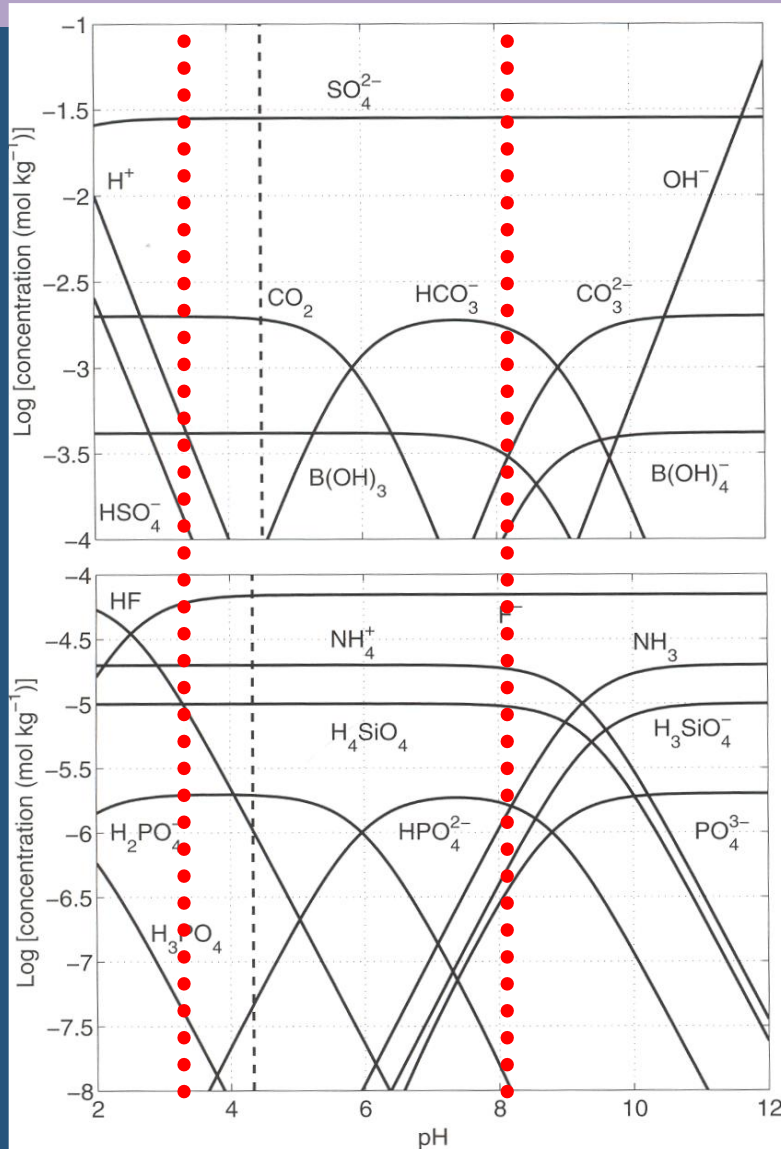
Pacífico Ocean



Total Alkalinity

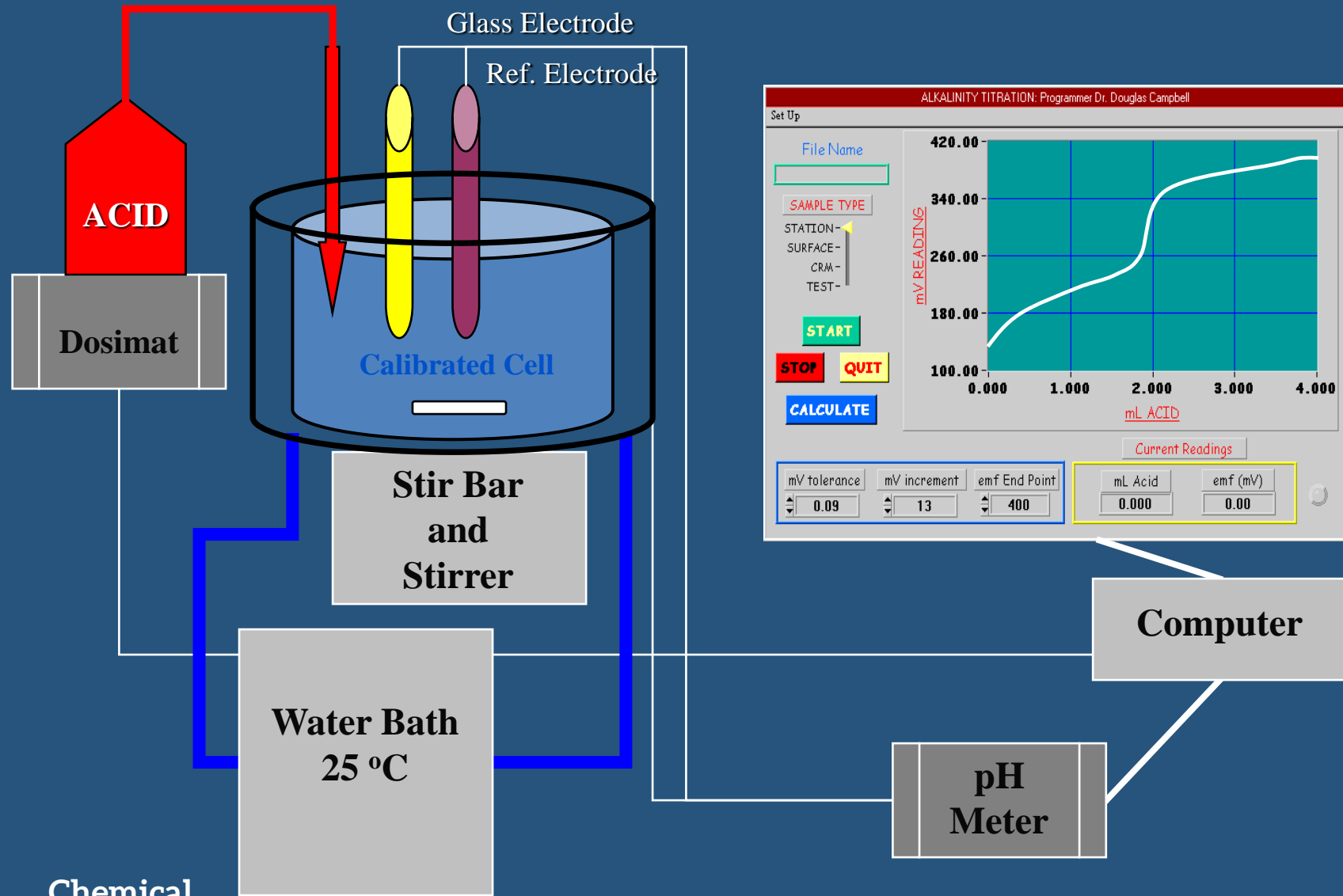


Alcalinidad Total



“The total alkalinity of natural waters is thus defined as the number of hydrogen ions equivalent to the excess of proton acceptors (bases formed from weak acids with a dissociation constant $K > 10^{-4.5}$, at 25°C and zero ionic strength) over proton donors (acids with $K > 10^{-4.5}$) in one kilogram of sample” *Dickson, 1994*

Total Alkalinity System



Total Alkalinity

$$\text{TA} = [\text{HCO}_3^-] + 2[\text{CO}_3^{2-}] + [\text{B(OH)}_4^-] + [\text{Nutrients}] + [\text{Extra-Bases}]$$

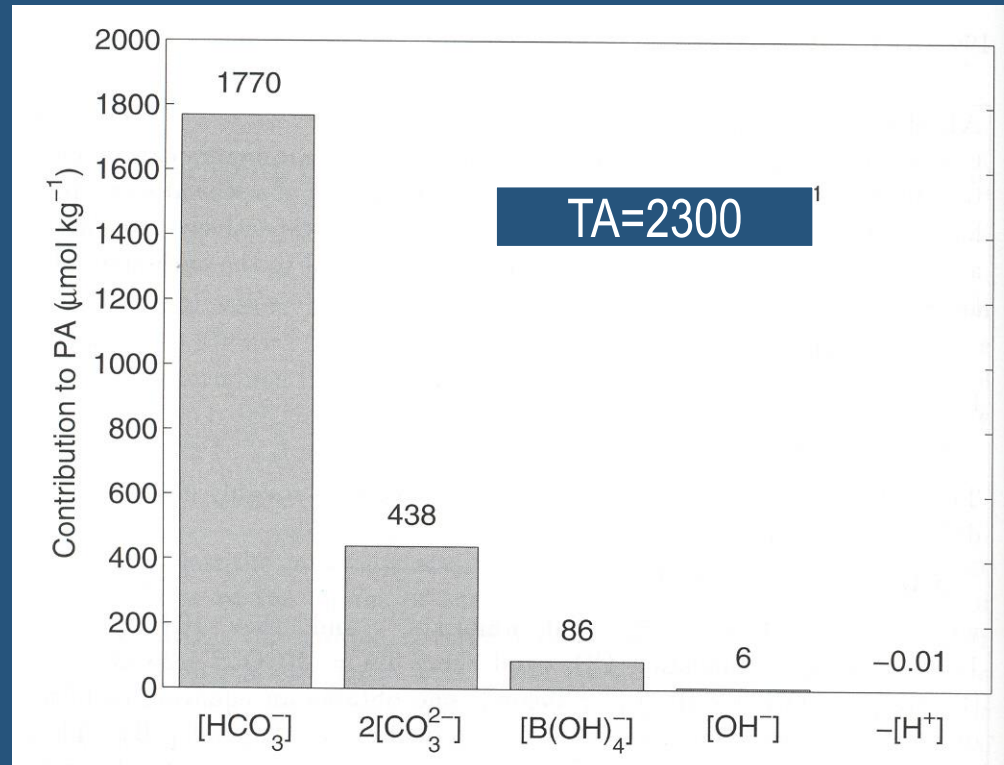
Known volume of seawater is acidified with a 0.1 N HCl solution to pH between 3.8 and 4.2.



Total Alkalinity

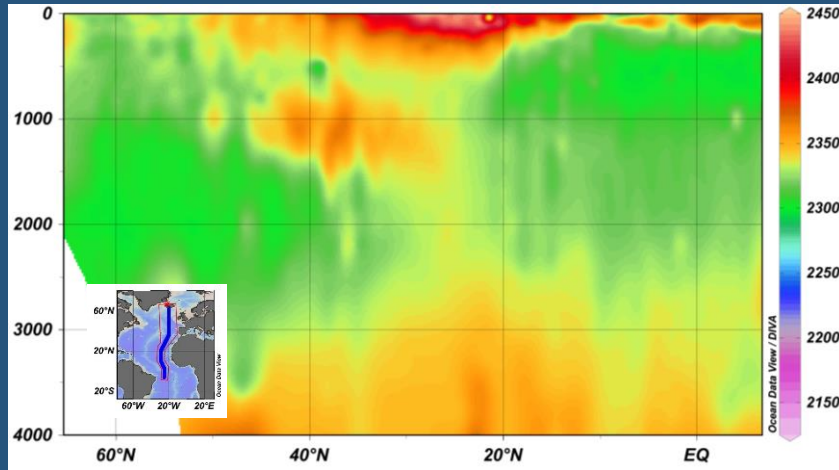
$$TA = [HCO_3^-] + 2 \cdot [CO_3^{2-}] + [B(OH)_4^-] + [OH^-] - [H_3O^+]$$

$$TA = [HCO_3^-] + 2 \cdot [CO_3^{2-}] \quad 96\% \text{ of TA}$$

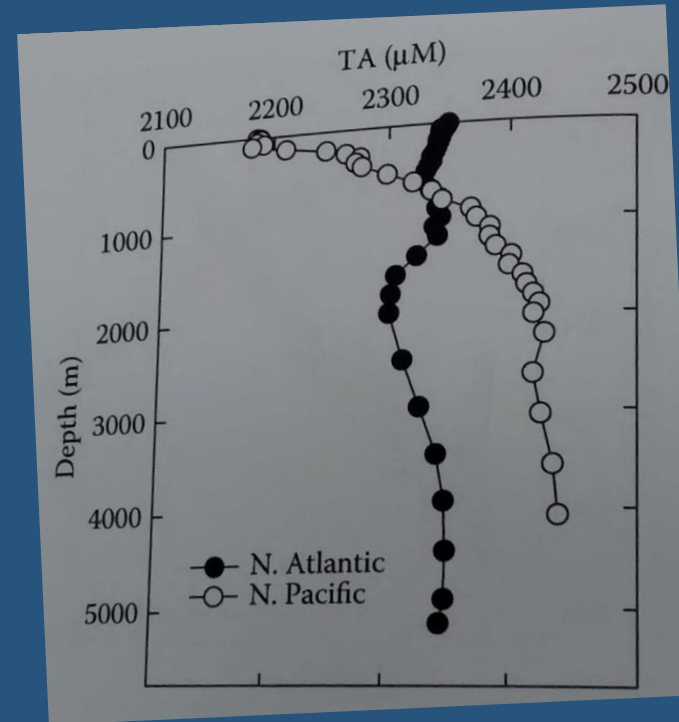
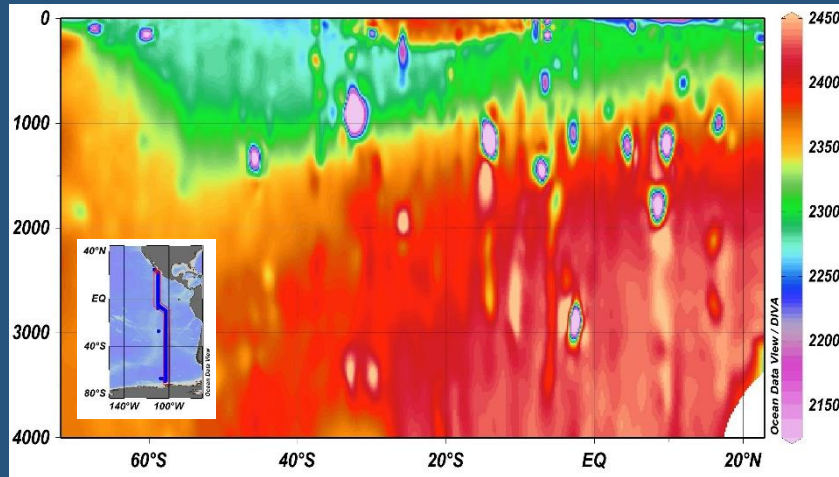


Total Alkalinity

Atlantic Ocean



Pacífico Ocean





CARBON DIOXIDE PARAMETERS IN SEA WATER

- Total Dissolved Inorganic Carbon

$$C_T = [\text{CO}_2] + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$

T, p independent

UNITS: moles per kilogram of solution (usually $\mu\text{mol kg}^{-1}$)

- Total Hydrogen Ion Concentration (pH)

$$\text{pH} = -\lg [\text{H}^+]$$

T, p dependent

UNITS: pH is dimensionless

But, total hydrogen ion concentration is in moles per kilogram of solution

- Partial Pressure of CO_2

(in air that is in equilibrium with the water sample)

$$p(\text{CO}_2) = x(\text{CO}_2) p = [\text{CO}_2] / K_0$$

T, p dependent

UNITS: pressure units (usually μatm)



TOTAL ALKALINITY

$$A_T = [\text{HCO}_3^-] + 2[\text{CO}_3^{2-}] + [\text{B(OH)}_4^-] + [\text{OH}^-] - [\text{H}^+]$$

T, p independent

UNITS: moles per kilogram of solution (usually $\mu\text{mol kg}^{-1}$)

Why are the concentrations of these species not affected by CO_2 dissolution in surface waters?

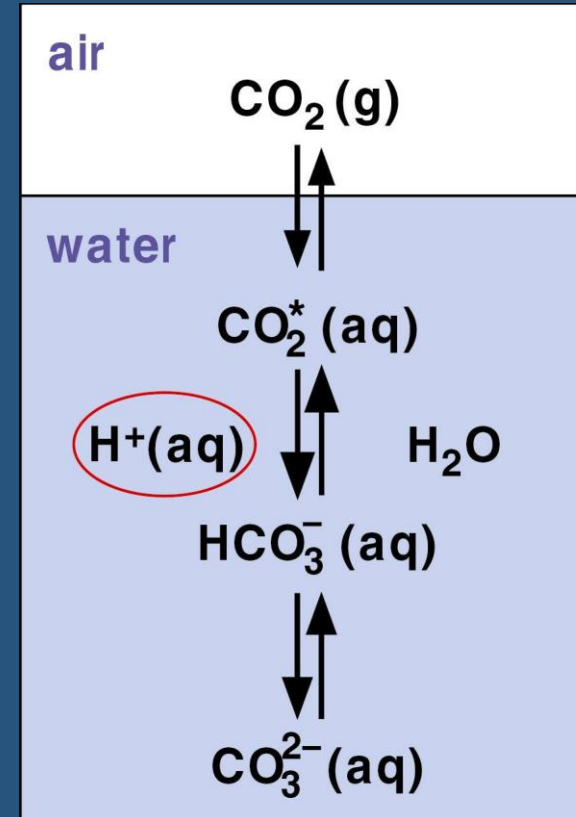
How does the Thermodynamics of CO₂ in seawater help us?

$$K_0 = [\text{CO}_2]/p(\text{CO}_2)$$
$$K_1 = [\text{H}^+][\text{HCO}_3^-]/[\text{CO}_2]$$
$$K_2 = [\text{H}^+][\text{CO}_3^{2-}]/[\text{HCO}_3^-]$$

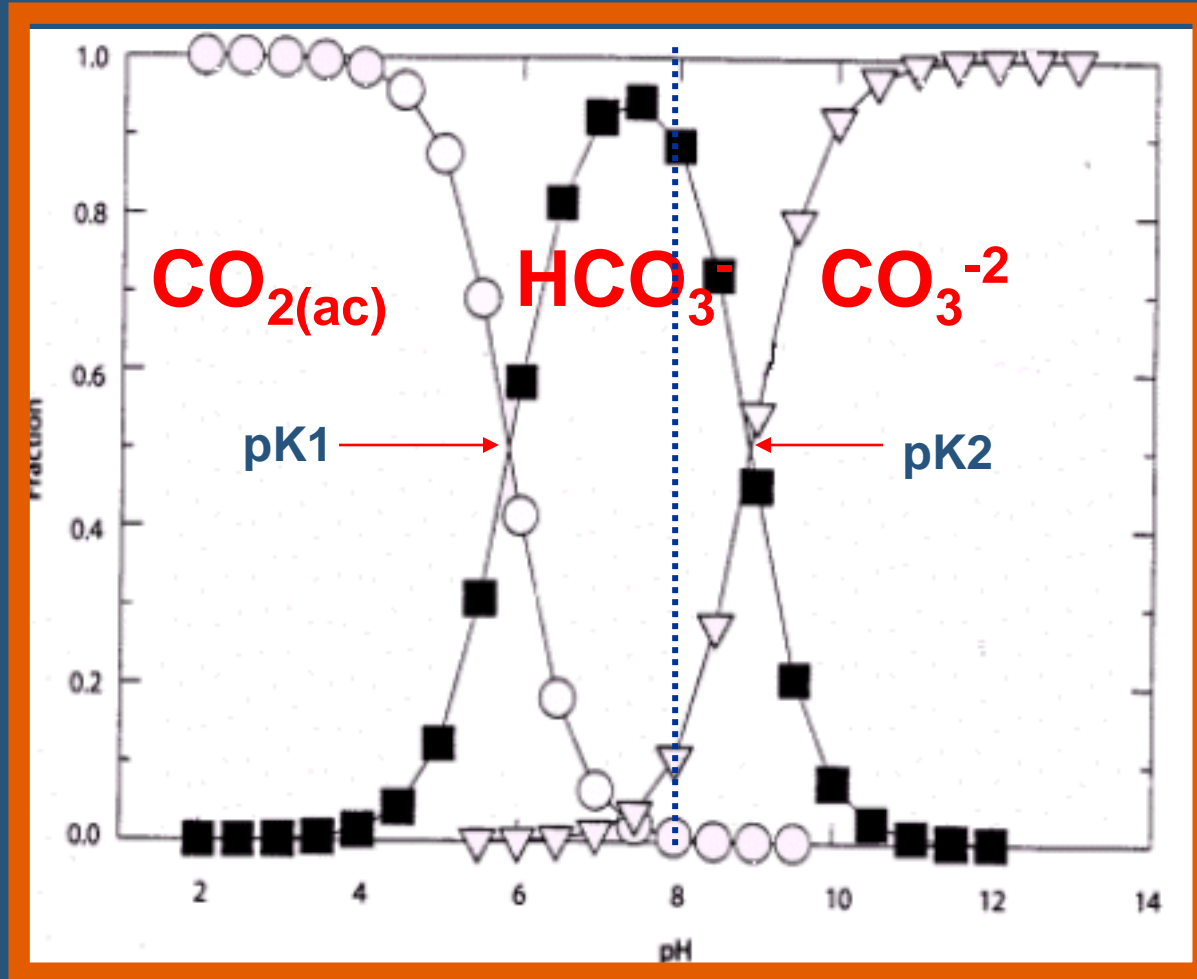
: T, S, P

CO₂ system parameters :

pH,
Total alkalinity (TA)
Dissolved inorganic carbon (DIC)
Partial Pressure of CO₂ (pCO₂)



“pK’s”



Millero et al., 1994

$$K_1 = \frac{[\text{H}^+][\text{HCO}_3^-]}{[\text{CO}_2]}$$
$$K_2 = \frac{[\text{H}^+][\text{CO}_3^{2-}]}{[\text{HCO}_3^-]}$$

“K” Values depend of S,
T, P.

Valores de K'

The values here are for S = 35, T = 25°C and P = 1 atm.

<u>Constant</u>	<u>Apparent Seawater Constant (K')</u>
K' _H	10 ^{-1.53}
K' ₁	10 ^{-6.00}
K' ₂	10 ^{-9.10}
K' _w	10 ^{-13.9}

Resume to this point:

- The CO₂ system in the seawater
- Components of the CO₂ system: pH, DIC, pCO₂ and alkalinity.
- At least to measure two of the four
- Use the CO2Sys program to estimate the different species from the carbon system
- “Next” process to control the different variables



Any pair of parameters describing the CO₂ system in a seawater sample can be used.

- Mathematically all combinations should be equivalent.
 - In practice this is not the case. Each of the terms has associated uncertainties. The uncertainties propagated through the calculations result in uncertainties in several calculated values.
 - In addition the uncertainties of the measured CO₂ system parameters have uncertainties in several of their equilibrium constants and in the total concentrations of the acid-base systems.
 - In addition the components of the alkalinity equation may be incomplete.
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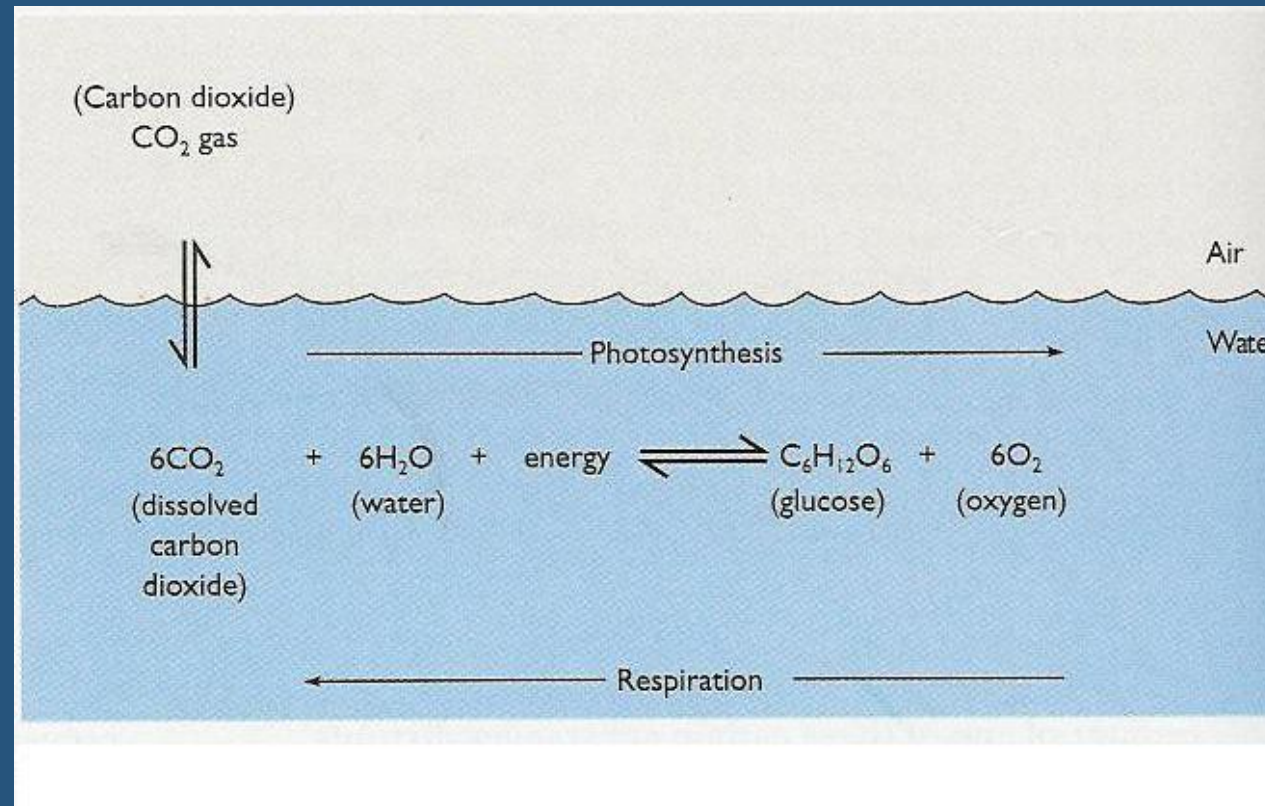
ESTIMATED UNCERTAINTIES FOR MEASUREMENT TECHNIQUES ON DISCRETE SAMPLES[†]

Parameter	State-of-the-art laboratory	State-of-the-art at-sea (suitable RMs)	Other techniques (suitable RMs)	Techniques not using RMs
Total alkalinity	1.2 $\mu\text{mol kg}$	2–3 $\mu\text{mol kg}$	4–10 $\mu\text{mol kg}$?
Total carbon	1.0 $\mu\text{mol kg}$	2–3 $\mu\text{mol kg}$	4–10 $\mu\text{mol kg}$?
pH	0.003	~0.005	0.01–0.02	?
p	1.0 μatm	~2 μatm	5–10 μatm	?

[†] Based on measuring surface oceanic CO₂ levels

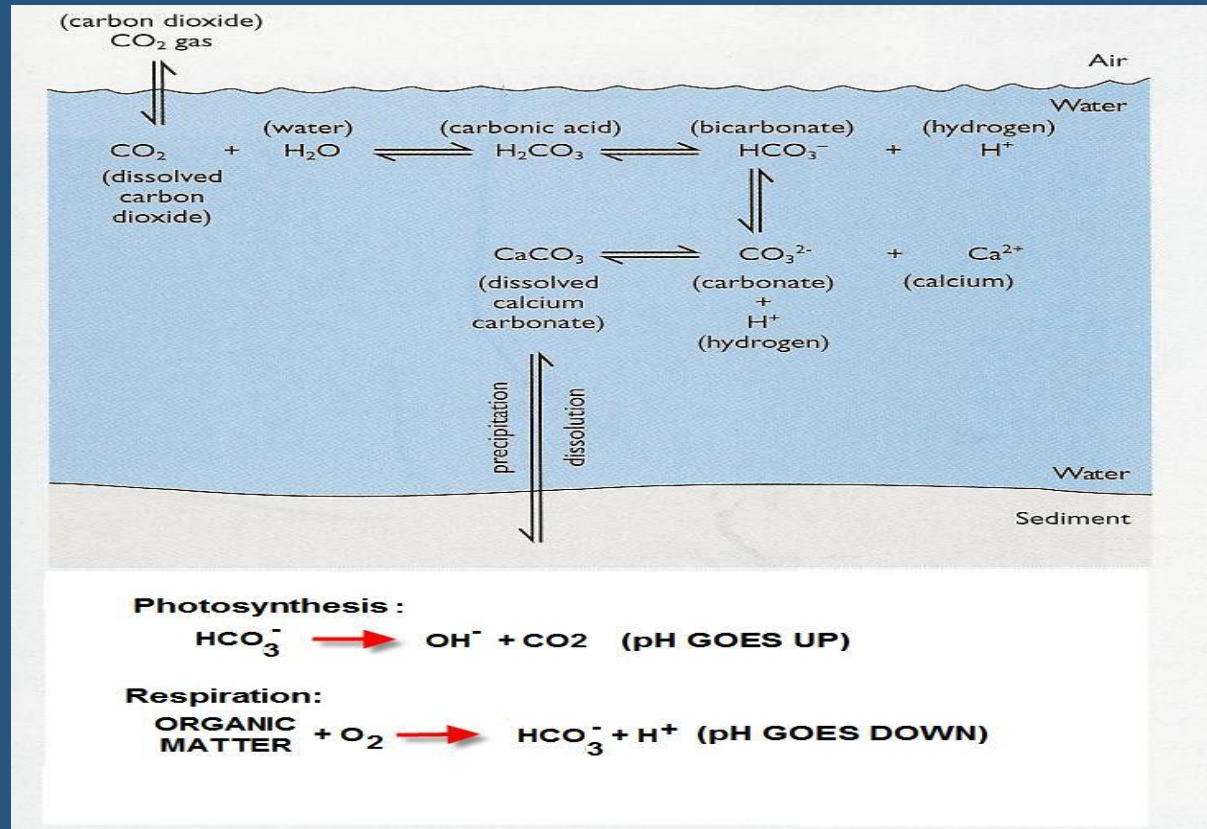
CO₂ FACTS

Photosynthesis ↓ use CO₂
Respiration increase ↑ CO₂



CO₂ FACTS

Photosynthesis change pH in the seawater



Controls on Ocean Distributions

A) **Fotosíntesis/Respiración**

Organic matter (approximated as CH_2O for this example) is produced and consumed as follows:



Then:



As CO_2 is produced during respiration we should observe:



The trends will be the opposite for photosynthesis.

B) **CaCO_3 dissolution/precipitation**



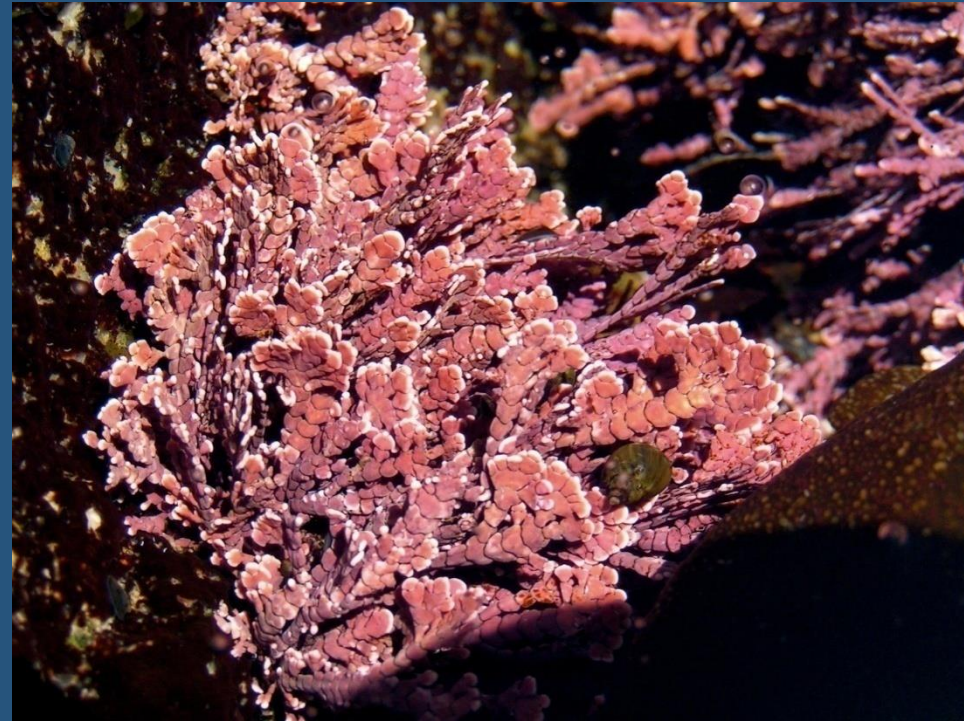
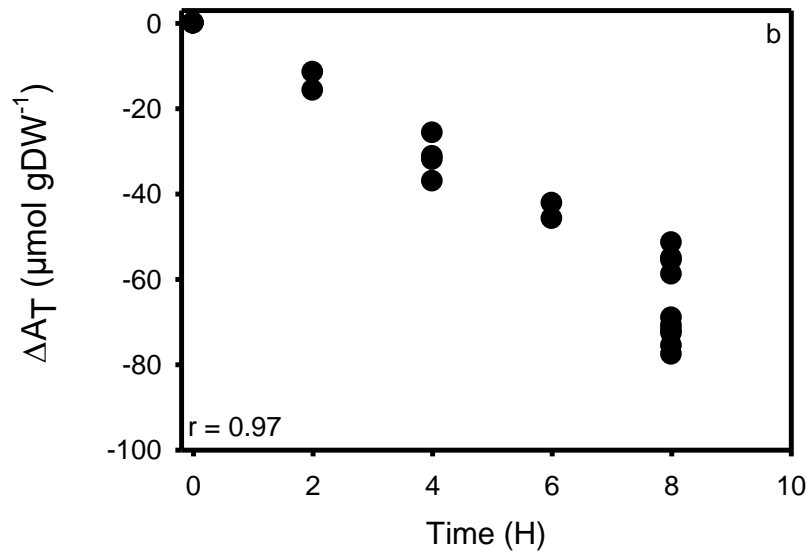
Also written as:



As $\text{CaCO}_3(\text{s})$ dissolves, CO_3^{2-} is added to solution. We should observe:



Calcification Rates and Acquisition of Inorganic Carbon In Geniculate Coralline Algae Distributed in the Coast Of Baja California Mexico



Changes in total alkalinity in 8h incubations with constant aeration. a, control method; filtered seawater with constant aeration; b, lineal decrement of total alkalinity normalized by dry weight in *Bossiella orbigniana*.

"The State of Saturation "The Magic Number"

$$\Omega_{\text{aragonite}} = [\text{Ca}^{+2}] [\text{CO}_3^{-2}] / K_{\text{ps arag}}$$

$$\Omega_{\text{calcite}} = [\text{Ca}^{+2}] [\text{CO}_3^{-2}] / K_{\text{ps cal}}$$

$\Omega \gg 1$ Organisms can build their shells "more easily"

$\Omega > 1$ Organisms can build shell

$\Omega < 1$ The shells "dissolve"

Adult Oyster Shell
—Formed from Calcite--Difficult to dissolve



Oyster larvae shell
—Formed from Aragonite—Easy to dissolve



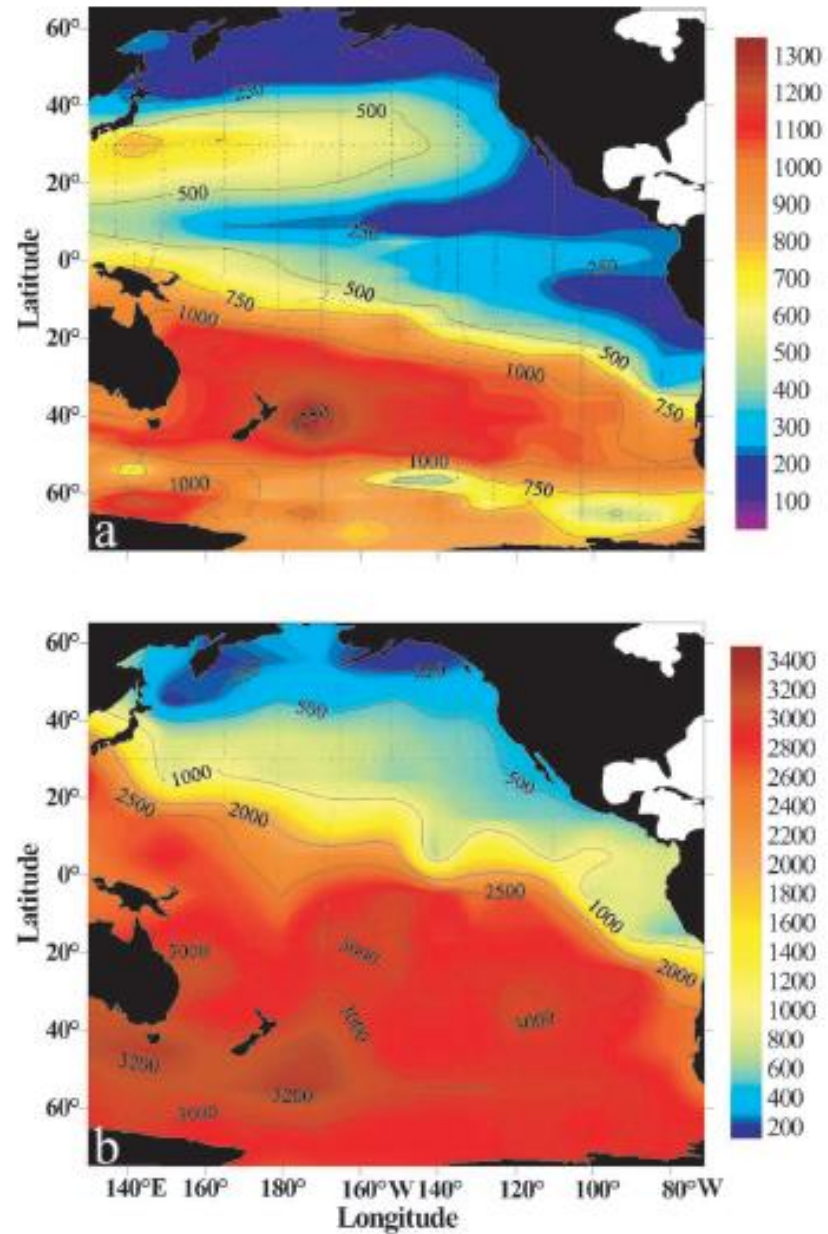


Figure 4. Saturation depth in meters for (a) aragonite and (b) calcite estimated from water column DIC and TA concentrations.

Aragonite



Both are CaCO_3

Calcite





To study the Ocean Acidification “OA” several “Tools” are required:

1. Understanding acid-base reactions in seawater
“Especially CO_2 ”
 2. Understand which CO_2 parameters (pH, Total Alkalinity, Dissolved Inorganic Carbon, pCO_2) are normally measured in seawater and how.
 3. **Have access to a computing tool to perform CO_2 system calculations such as CO2Sys or CO_2 Cal.**
 4. Have access to analytical equipment (and training) to perform measurements of carbon variables.
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